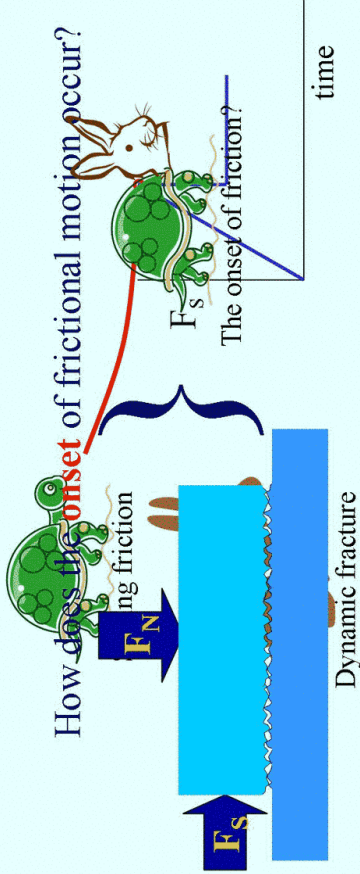


Detachment fronts and the onset of friction

Shmuel M. Rubinstein, Gil Cohen, and Jay Fineberg

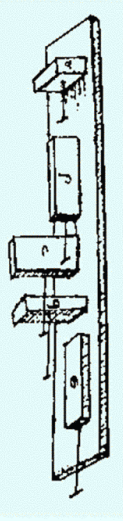
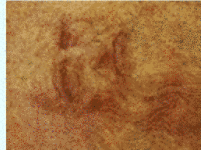
The Racah Institute of Physics, The Hebrew University of Israel
Jerusalem 91904, Israel



Reference: S. M. Rubinstein, G. Cohen, and J.F., Nature **430**, 1005-1009 (2004).

The Study of Friction

Leonardo Da Vinci (1452-1519)

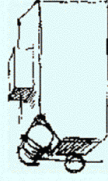


1. The areas in contact have no effect on friction.
2. If the load of an object is doubled, its friction will also be doubled.
→ $F_s \propto F_N$
3. Friction is related to the roughness of the material in question



Guillaume Amontons (1663-1705)

Charles August Coulomb (1736-1806)

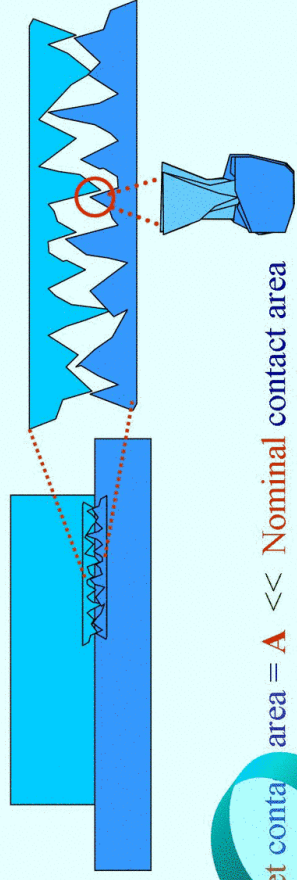


→ “Static” and “Dynamic” friction:

$$F_s = \mu_s F_N \quad (v = 0)$$

$$F_s = \mu_D F_N \quad (v > 0)$$

Explaining the Amontons-Coulomb law:



- Net contact area = $A \ll \ll$ Nominal contact area
- Huge pressures at the contact points deform the materials
- A grows until pressure = yield strength, σ_Y
 $\rightarrow A = F_N / \sigma_Y$
- Motion: Fracture of contacts
 $\rightarrow F_S = \tau_C A$, where τ_C = shear strength

$$\rightarrow F_S = \tau_C A = F_N (\tau_C / \sigma_Y) = \mu F_N$$

F. Philip Bowden and David Tabor (1950)

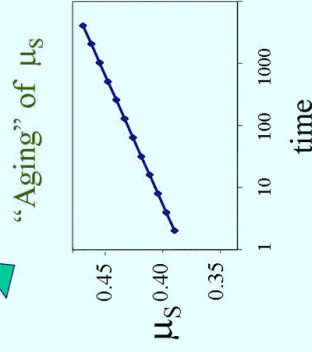
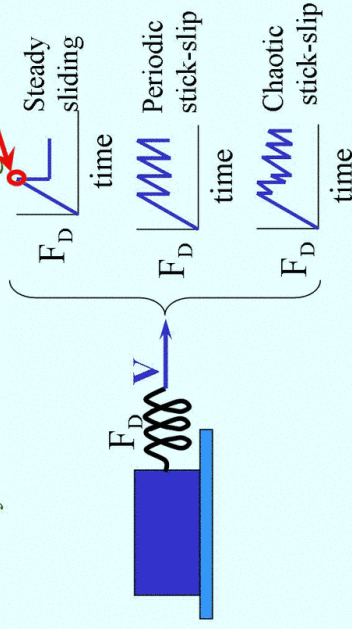
Sliding friction:

Rate and State Friction Laws

$$F_S = [\mu_0 + a \cdot \log(V/V_0) + b \cdot \log(V_0 \theta / L)] \cdot F_N$$

$$d\theta/dt = 1 - \theta V / L \quad (\theta \text{ is a "state variable" and } V \text{ the sliding velocity})$$

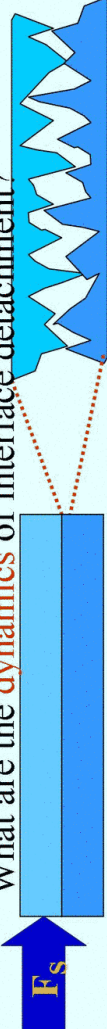
Dynamics of frictional sliding



Dieterich, J., *Pure Appl. Geophys.* **116**, 790 (1978)
 Ruina, A.L., *J. Geophys. Res.* **88**, 10359 (1983)
 Baumberger, T., Berthoud, P. & Caroli, *Phys. Rev.* **B60**, 3928 (1999)

Detachment of the Interface ⇔ Frictional motion

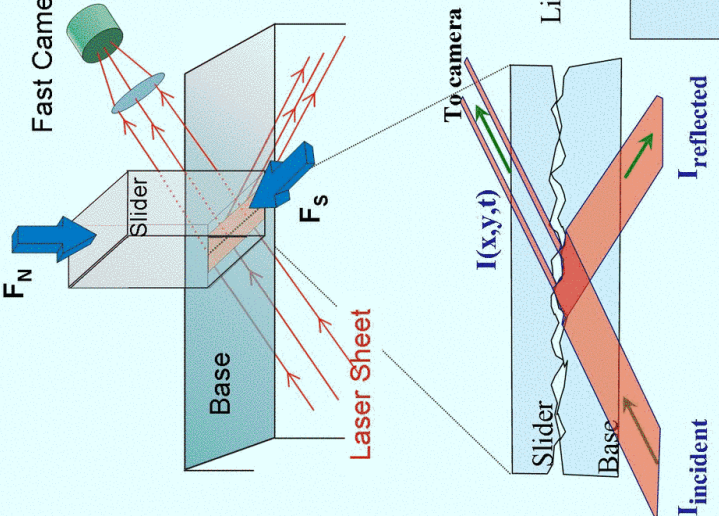
What are the **dynamics** of interface detachment?



- Statics: **Net** contact area = $A \ll \ll$ **Nominal** contact area
- Dynamics: **Fracture** of the contacts occurs when $F_s = \tau_c A$, where τ_c = material shear strength.

F. P. Bowden and D. Tabor (1950)

Our experiment: **What are the dynamics of interface detachment??**



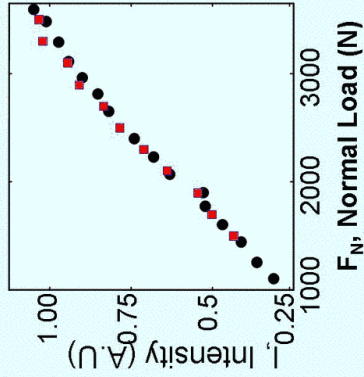
Mechanical details:
 Base ⇔ plexiglass or glass
 Slider ⇔ plexiglass
 Interface roughness: $\sim 1 \mu\text{m}$
 Normal stresses: 1-6 MPa
 Interface dimensions: 5.5 x 150 mm
 Camera:
 Frame size = $1280 \times N$ pixels
 Frame rate = $500,000/(N+1)$ frames/sec

Incident angle adjusted for
Total Internal Reflection at interface

Light is only transmitted at the **actual points of contact**

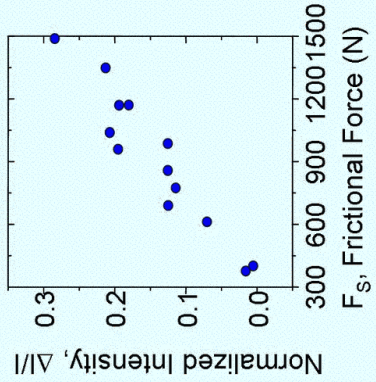
$I(x,y,t)$ = **Instantaneous net contact area**
 at every location x,y and time t **along the entire interface**

Does the method work? “Static” measurements:



The net contact area increases linearly with F_N
 $I \propto F_N \rightarrow A \propto F_N$
 Dieterich, J. H. & Kilgore, B. D. *Pure and Applied Geophysics* 143, 283-302 (1994).

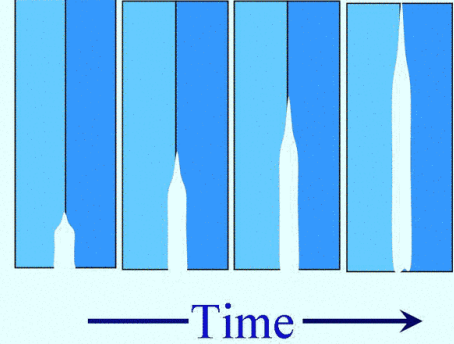
I is a **direct** measure of both the net contact area, A , and the frictional force, F_S
 $F_S = \mu F_N \propto A \propto I$



Static measurements \Leftrightarrow **Verification of Bowden and Tabor!**

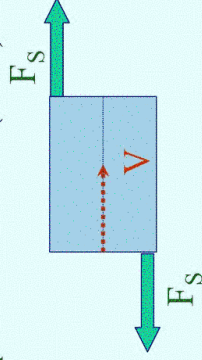
Detachment of the Interface \Leftrightarrow Frictional motion

What are the **dynamics** of interface detachment?



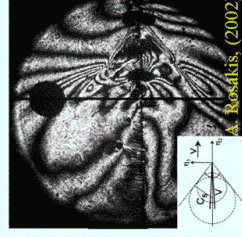
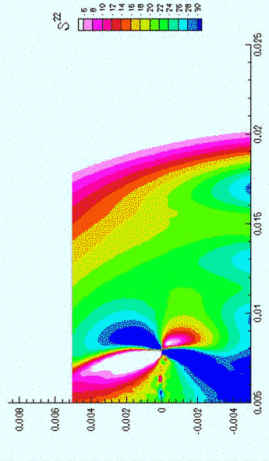
Is interface detachment governed by the propagation of coherent (**crack-like**) fronts?
 Is interface detachment a **Dynamic Fracture Problem**?

Friction as a fracture process: Mode II (Shear) Fracture



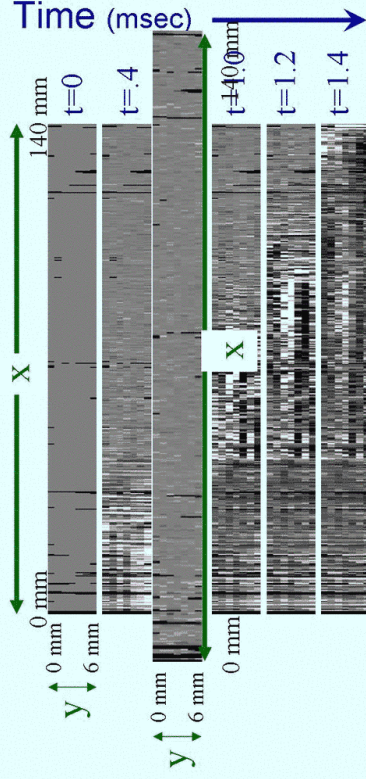
Shear fracture:

- Can only occur along a **weak interface** – otherwise tensile fracture occurs
- Shear cracks with **point-like stress field singularities** at their tip can propagate for all $V < V_R$ and for the special velocity of $V = \sqrt{2}V_S$ (V_S is the shear wave speed).
- Any other shear fracture velocities can only occur for a **spatially distributed dissipation zone**^{2,3} ahead of the fracture.



1. Freund, L. B., *J. Geophys. Res.* **84**, 2199-2209 (1979)
2. Rosakis, A. J., Samudrala, O. & Coker, D. *Science* **284**, 1337-1340 (1999); Needleman, A., 1999, *J. Mech. Phys. Solids*, **66**, 847. K.
3. Xia, A. J., Rosakis, and H. Kanamori, *Science* **303**, 5665 (2004).

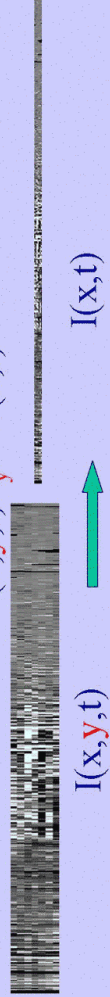
The Dynamic behavior of the interface: $I(x,y,t) = ?$



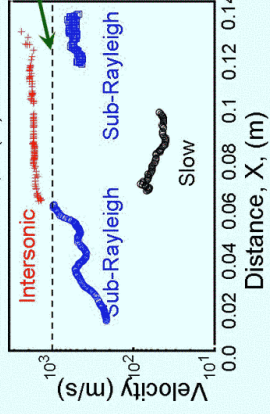
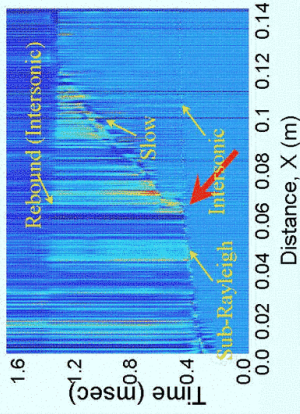
Interface detachment occurs by means of **crack-like fronts** propagating in the x direction along the interface

As the fronts are quasi-1D, we will now (for simplicity) consider:

$$\langle I(x,y,t) \rangle_y = I(x,t)$$



Properties of detachment fronts:

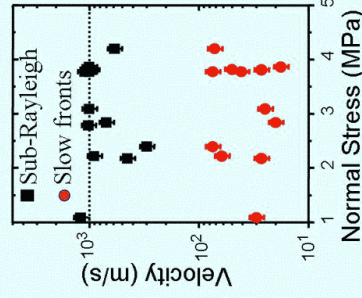


Dynamics:

- Both Intersonic and Slow fronts are emitted upon the arrest of Sub-Rayleigh fronts at $V \sim V_R$

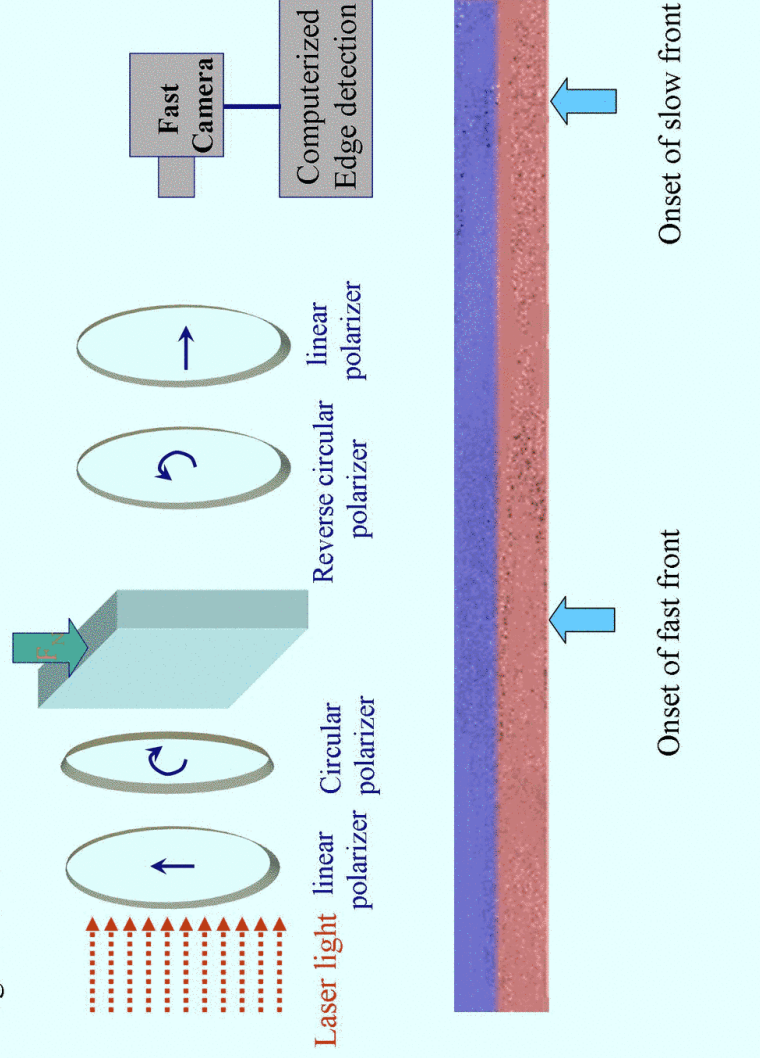
Three distinct types of detachment fronts exist:

- Intersonic¹ fronts $1.7V_R > V > 1.3V_R$
- Sub-Rayleigh fronts $1.0V_R > V > 0.3V_R$
- “Slow” Detachment fronts $0.1V_R > V > 0.02V_R$



¹ Rosakis, A. J., Samudrala, O. & Coker, D. *Science* **284**, 1337-1340 (1999).
 Xia, K., Rosakis, A. J. & Kanamori, H.. *Science* **303**, 1859-1861 (2004).

A photoelastic measurement of the onset and movement of a slow front moving from right to left.



Experiments in Granite

M. Ohnaka and L. Shen, JGR 104, 817 (1999)

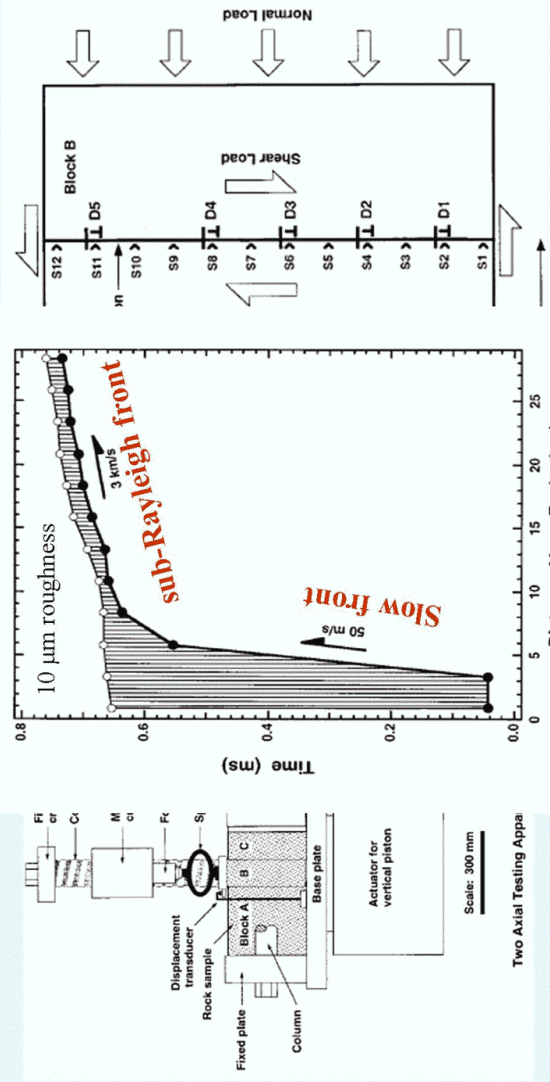
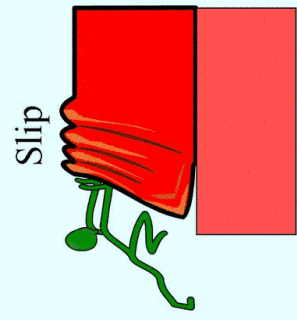


Figure 26. Space-time view of the nucleation zone for a slip failure event (E40117comb) on the extremely smooth fault. The hatched portion indicates the zone in which the slip weakening is proceeding with time.

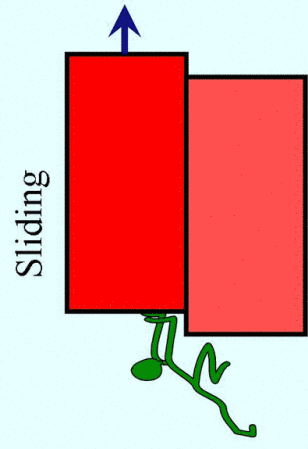
Frictional slip vs. Frictional Sliding



Local displacement along the interface

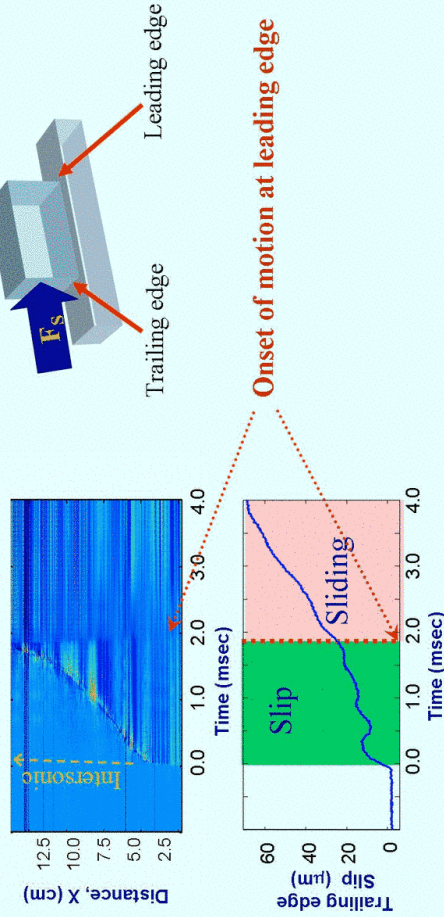
Interface Fracture

The transition from slip to sliding?
The transition from static to dynamic friction?



Global translation of the interface

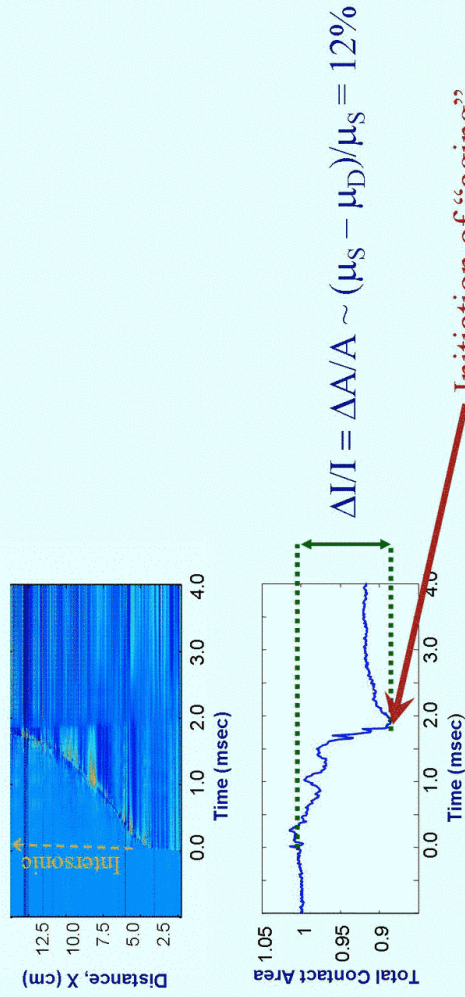
Slip along the interface:



- No slip occurs upon passage of **Intersonic** fronts
- Slip at the **trailing edge** occurs during propagation of both **Sub-Rayleigh** and **Slow Detachment** fronts
- No slip at the **leading edge** occurs until the arrival of either the **Slow** or **Sub-Rayleigh** fronts – at that time **sliding** of the blocks occurs

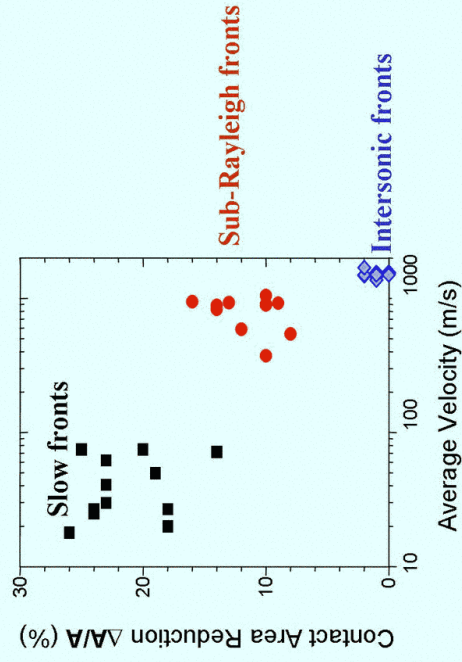
Contact area reduction by detachment fronts:

The transition from static to dynamic friction



- The total reduction of surface area by the slow and sub-Rayleigh fronts drives the transition from **“static”** to **“dynamic”** friction.
- Aging of the contact surface is initiated **immediately upon the onset of sliding**

Efficiency of the different fronts in contact area reduction:



- Intersonic fronts produce a **negligible** reduction of contact area
- Slow detachment fronts are **twice as efficient** as Sub-Rayleigh fronts in reducing the net contact area

Summary: The transition from static to dynamic friction
or
the dynamics of fault nucleation

1. Slip occurs via **three different types of detachment fronts** which propagate at highly different velocities.
2. Both **slow** detachment and **intersonic** fronts are emanated upon **arrest of sub-Rayleigh fronts** – upon their arrival at V_R
3. All three detachment fronts **nearly always** are **observed** – either alone or together with the other fronts
4. **Negligible slip** and **contact area reduction** result from the passage of **intersonic** fronts.
5. **Slow detachment fronts** are twice as **efficient** as sub-Rayleigh fronts in **contact area reduction**

Relevance to earthquakes: The dynamics of fault nucleation

Sub-Rayleigh fronts ⇔ “Standard” earthquakes ($0.2V_R < V < 0.9V_R$)
 Intersonic fronts ⇔ “Intersonic” earthquakes (e.g. Izmit 1999¹)
 Slow detachment fronts ⇔ ???

Slow earthquakes² = slow detachment fronts?

Characteristics of “slow” or “silent” earthquakes

- Significant slip/strain release (measured with GPS /strainmeters)
- “silent” – having a weak atypical seismic signature.

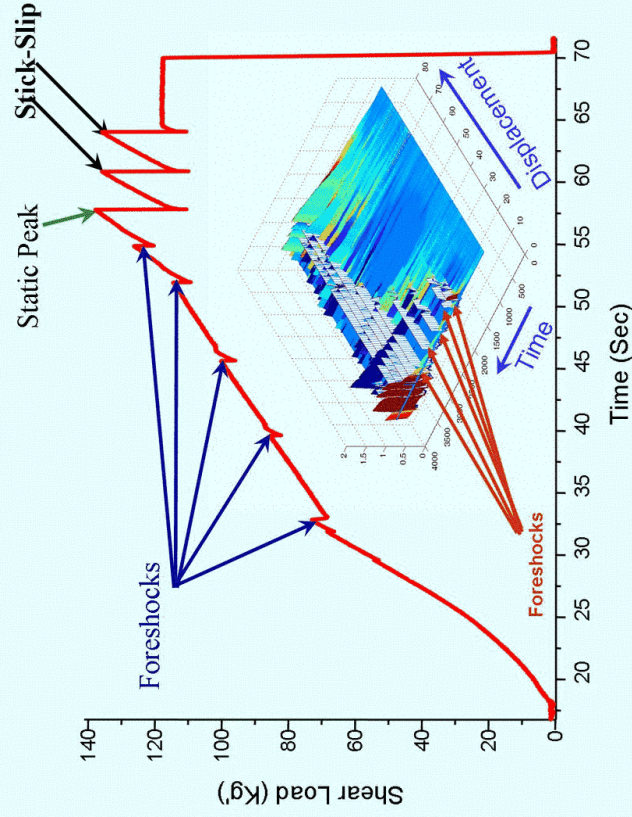
Question:

Slow detachment fronts are nearly always observed in experiments.
 Are silent earthquakes more common than currently believed?

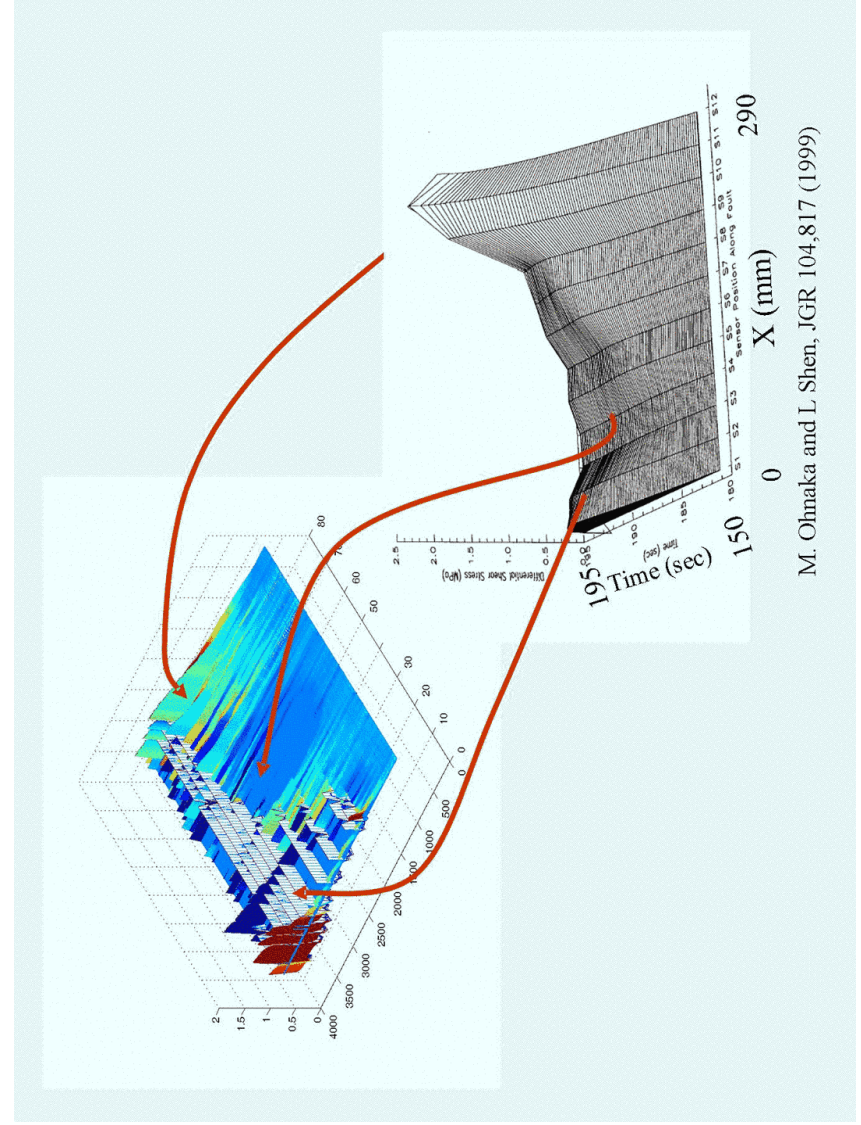
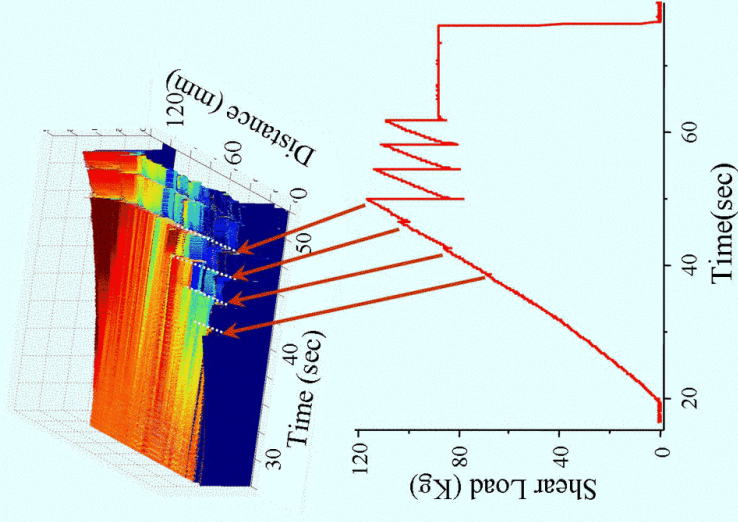
¹ Bouchon, M. et al. *Geophys. Res. Lett.* **28**, 2723-2726 (2001).

² Crescentini, L., Amoruso, A. & Scarpa, R. *Science* **286**, 2132-2134 (1999);
 Linde, A. T. & Sacks, I. S. *Earth and Planetary Science Letters* **203**, 265-275 (2002).
 Rogers, G. & Dragert, H. *Science* **300**, 1942-1943 (2003).

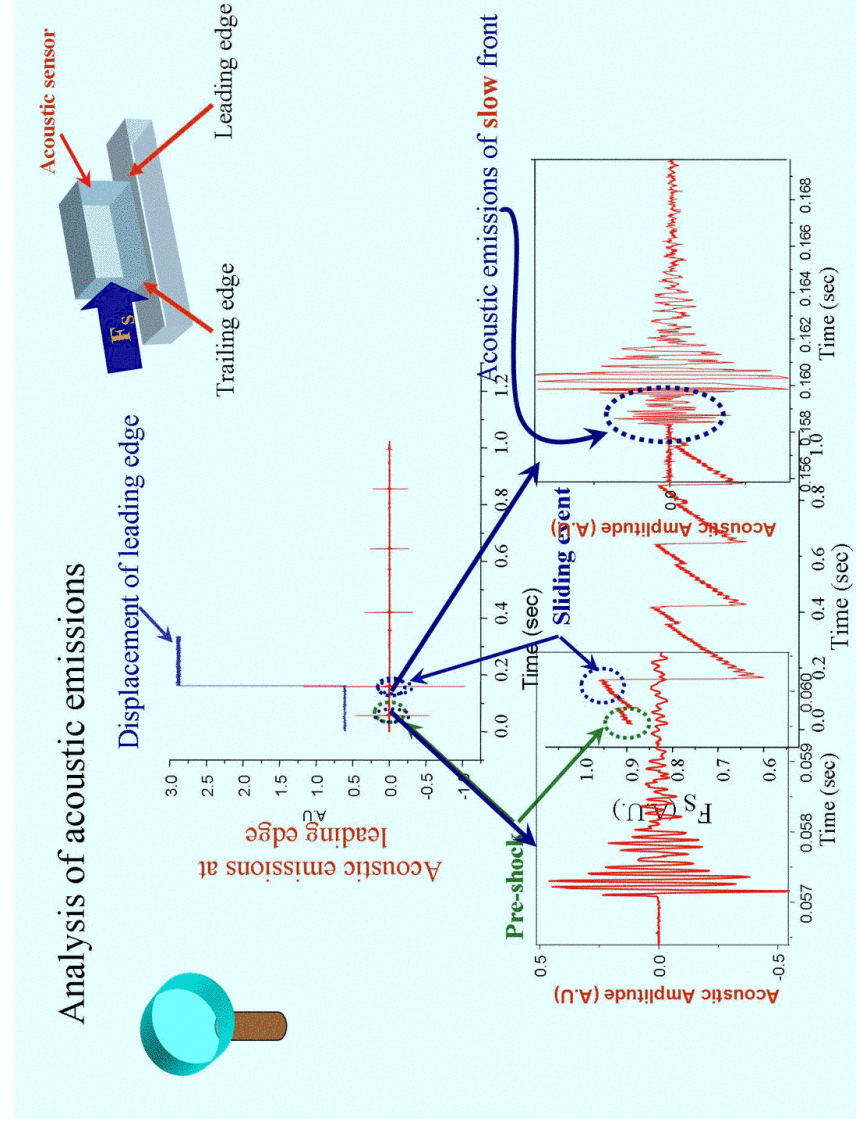
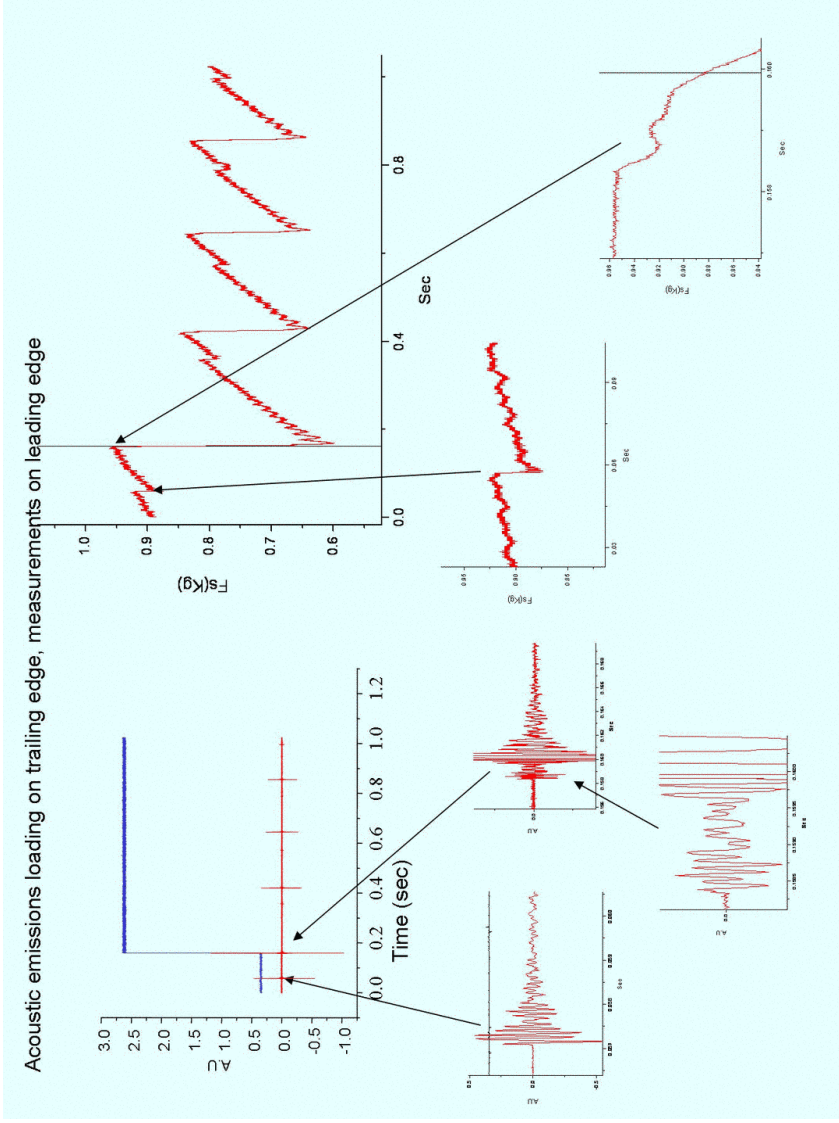
Foreshocks and the loading history



The foreshocks change the normal stress distribution along the interface



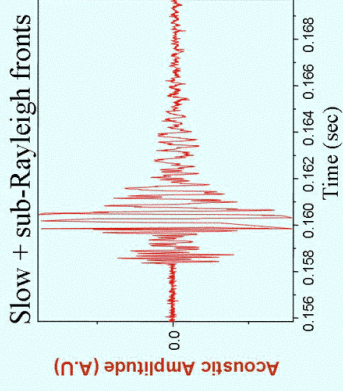
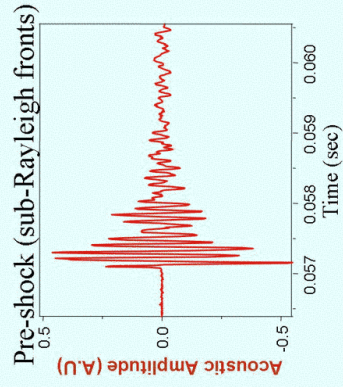
M. Ohnaka and L. Shen, JGR 104,817 (1999)



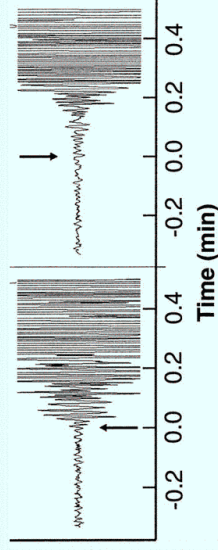
Slow earthquakes = slow detachment fronts?

Characteristics of “slow” or “silent” earthquakes

- Significant slip/strain release (measured with GPS /strainmeters)
- “silent” – having a weak atypical seismic signature.



The acoustic (“seismic”) emission amplitudes of the slow fronts are an order of magnitude **smaller** than those of the sub-Rayleigh fronts



Seismic signal from the magnitude 7

1994 Romanche transform earthquake

McGuire, J. J., Ihmlé, P. F. & Jordan, T. H.,

Science 274, 82–85 (1996).