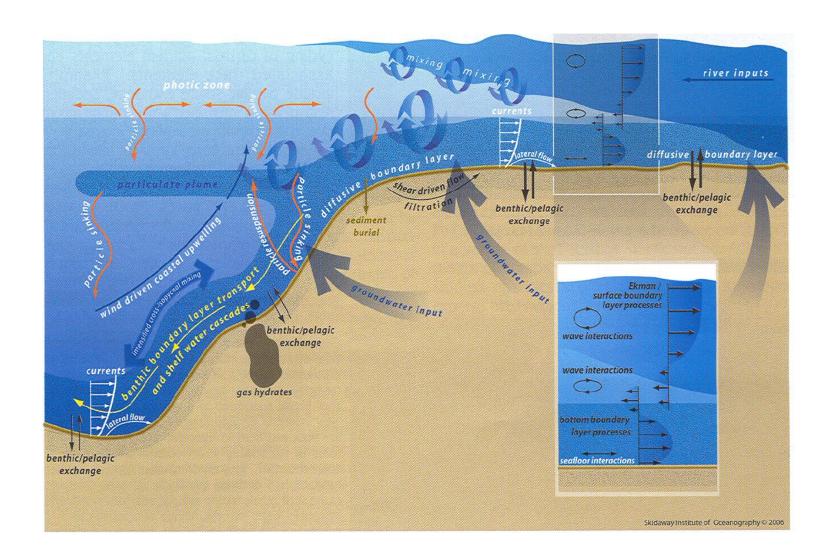
Modeling of Particle Flows in Aquatic Environments

Eckart Meiburg UC Santa Barbara

- Continuum approach
 - 'single-fluid' approach
 - extension from lab scales to field scales
- Grain-scale approach
 - erosion, deposition, segregation
 - concentration-dependent rheology / upscaling
- Summary
- Outlook



Coastal margin processes



Framework: Dilute flows

Assumptions:

- volume fraction of particles $< O(10^{-2} 10^{-3})$
- particle radius « particle separation
- small particles with negligible inertia

Dynamics:

- effects of particles on fluid continuity equation negligible
- coupling of fluid and particle motion primarily through momentum exchange, not through volumetric effects
- particle loading modifies effective fluid density
- particles follow fluid motion, with superimposed settling velocity
- → 'single-fluid' approach

Moderately dilute flows: Two-way coupling (cont'd)

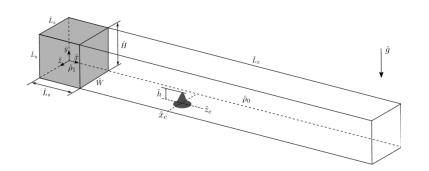
$$\nabla \cdot \vec{u}_f = 0$$

$$\frac{\partial \vec{u}_f}{\partial t} + \left(\vec{u}_f \cdot \nabla\right) \vec{u}_f = -\nabla p + \frac{1}{Re} \nabla^2 \vec{u}_f + c \, \vec{e}_g$$
 effective density
$$\frac{\partial c}{\partial t} + \left[\left(\vec{u}_f + \vec{U}_s\right) \nabla\right] c = \frac{1}{Sc \, Re} \nabla^2 c$$
 settling velocity

$$Re = \frac{u_b L}{\nu}$$
 , $Sc = \frac{\nu}{D}$, $U_s = \frac{u_s}{u_b}$

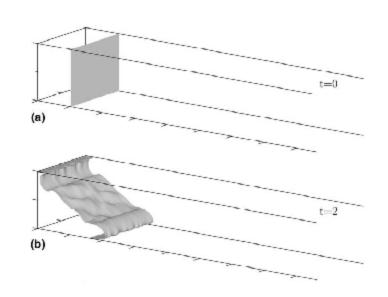
Model problem (with M. Nasr-Azadani)

Lock exchange configuration



Dense front propagates along bottom wall

Light front propagates along top wall



Computational approach for flow over complex geometry

- second order central differencing for viscous terms
- third order ENO scheme for convective terms
- third order TVD Runge-Kutta time stepping
- projection method to enforce incompressibility
- domain decomposition, MPI
- employ PETSc (developed by Argonne Nat'l Labs) package
- non-uniform grids
- immersed boundary method for complex bottom topography

Lock exchange configuration

Flow of turbidity current around localized seamount

Entry #: 84228

Particle-laden currents interacting with complex bottom topography: a numerical investigation

Mohamad M. Nasr-Azadani and Eckart Meiburg

University of California Santa Barbara

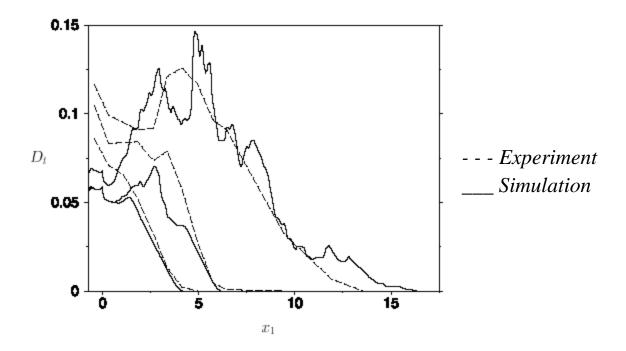
- turbidity current develops lobe-and-cleft instability of the front
- current dynamics and depositional behavior are strongly affected by bottom topography

$$Re_{sim} = 2{,}000: u_b \approx 2cm/s, L \approx 10cm, \nu \approx 10^{-6}m^2/s$$

→ simulation corresponds to a laboratory scale current, not field scale!

Deposit profiles

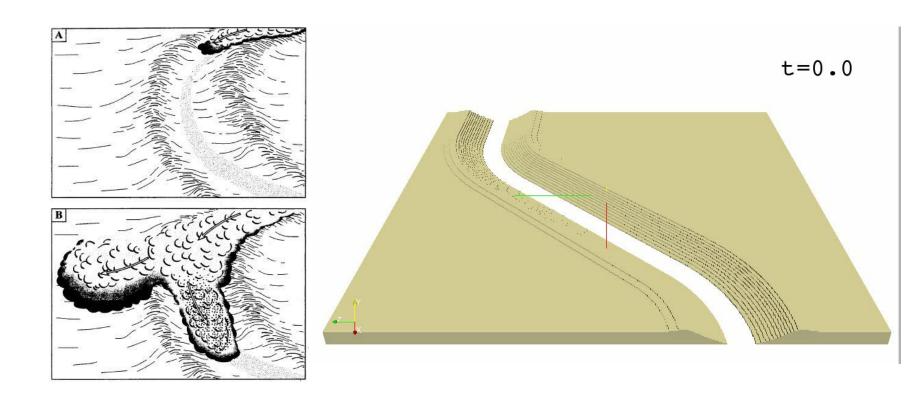
Comparison of transient deposit profiles with experimental data of de Rooij and Dalziel (1998)



• simulation reproduces experimentally observed sediment accumulation

Turbidity current/sediment bed interaction

'Flow stripping' in channel turns: lateral overflows

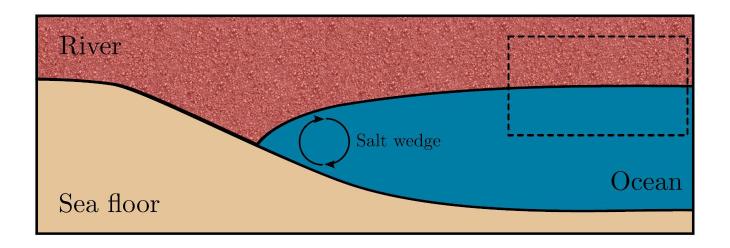


Sedimentation from river plumes (w. P. Burns)

Hypopycnal river plumes:

density of the river (fresh water + sediment) < density of ocean (water + salinity)

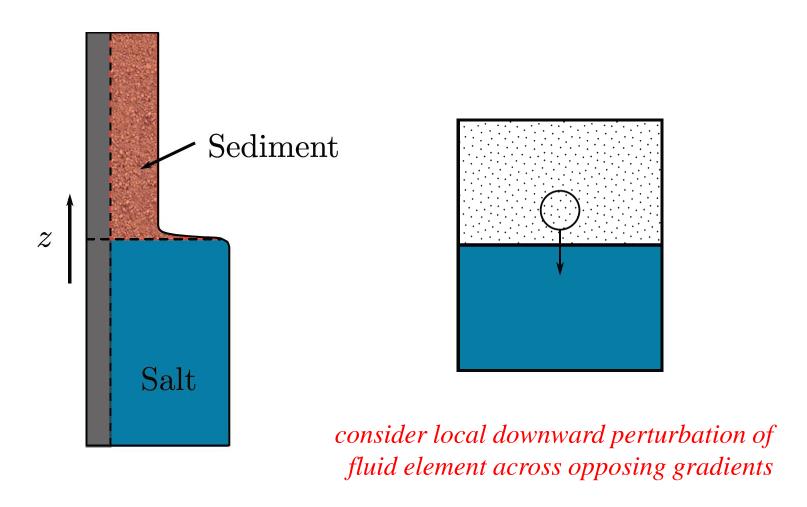
→ river outflow propagates along the ocean surface



• focus on the downstream density stratification

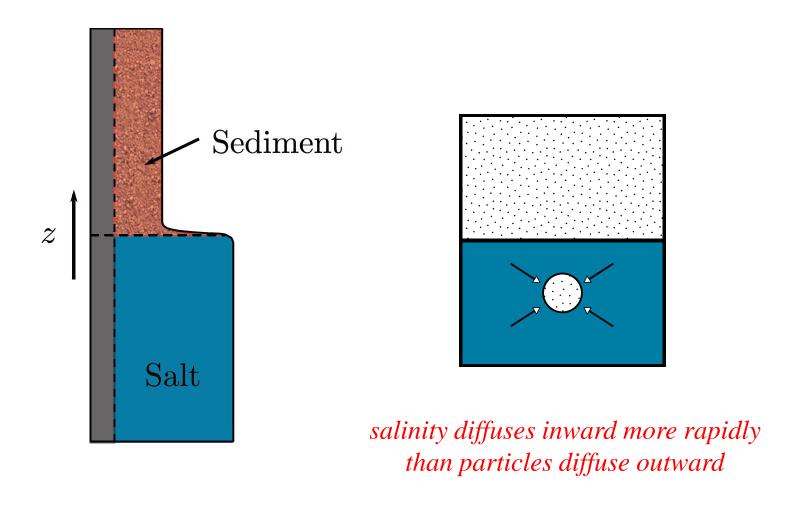
Sedimentation from river plumes: Double-diffusion

Base density profile:



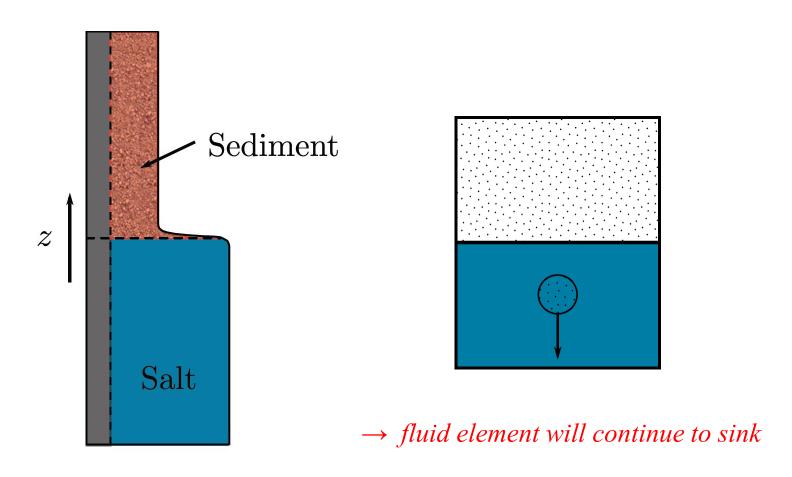
Sedimentation from river plumes: Double-diffusion

Base density profile:

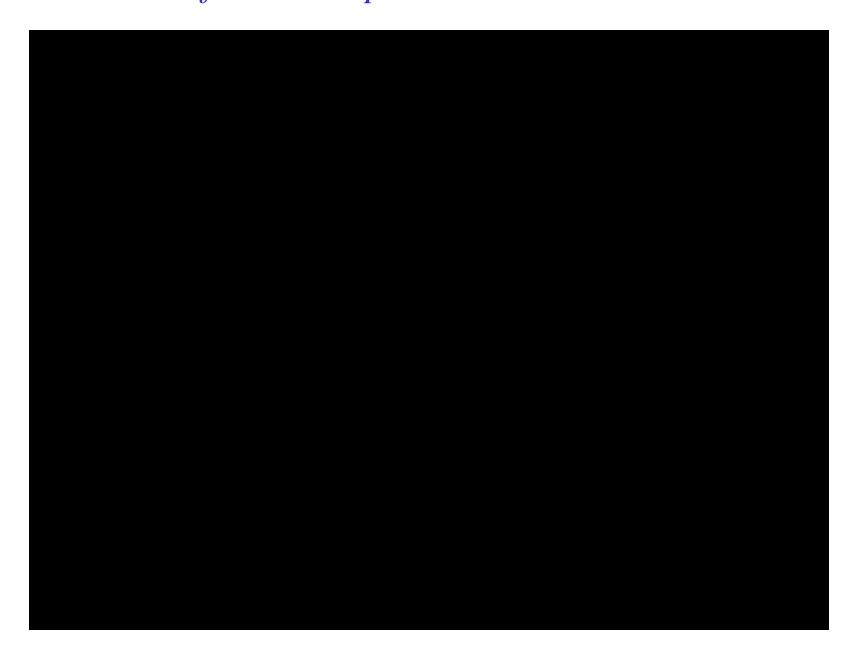


Sedimentation from river plumes: Double-diffusion

Base density profile:



Sedimentation from river plumes: Numerical simulations



DNS simulations

Strengths:

- accurately reproduce continuum-scale physics
- provide very detailed information on mixing, dissipation etc.
- require a minimum of empirical modeling assumptions

Current challenges:

- computationally very expensive
- limited to small Reynolds numbers, laboratory scale currents
- limited to dilute, depositional currents carrying small particles
- difficult to correctly capture erosion, bedload transport
- no direct particle/particle interactions (collisions)
- limited ability to reproduce segregation of different particle sizes
- no hindered settling, no concentration-dependent rheology

Alternative approach:

• two-fluid models: wider applicability, but require empirical closures

Why can we not do a DNS simulation at $Re=10^9$?

- Re is a measure of the ratio of the largest ("integral") length scale L of the flow to the smallest ("Kolmogorov") length scale η, at which kinetic energy is dissipated into heat
- turbulence theory shows that $\frac{L}{\eta} = Re^{3/4}$
- DNS, which resolves all scales, needs to have grid spacing $\Delta x \sim \eta$, and computational domain size $\sim L \rightarrow number$ of grid points in each direction $N \sim Re^{3/4}$. For 3D simulation $N_x \cdot N_y \cdot N_z \sim Re^{9/4}$. Time step $\Delta t \sim \Delta x \rightarrow N_x \cdot N_y \cdot N_z \sim Re^{9/4}$.

Computational effort $E \sim N_x \cdot N_y \cdot N_z \cdot \Delta t^{-1} \sim Re^3!!$

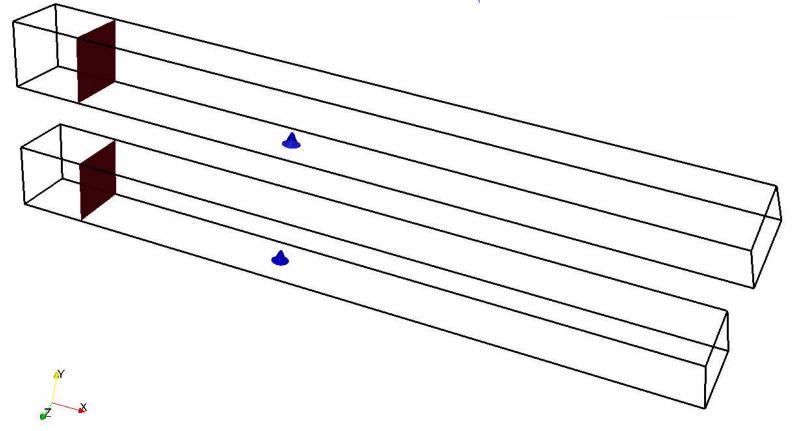
• field scale simulation would require 10¹⁸ times effort of lab scale simulation

How can we perform simulations at field scale?

Key idea:

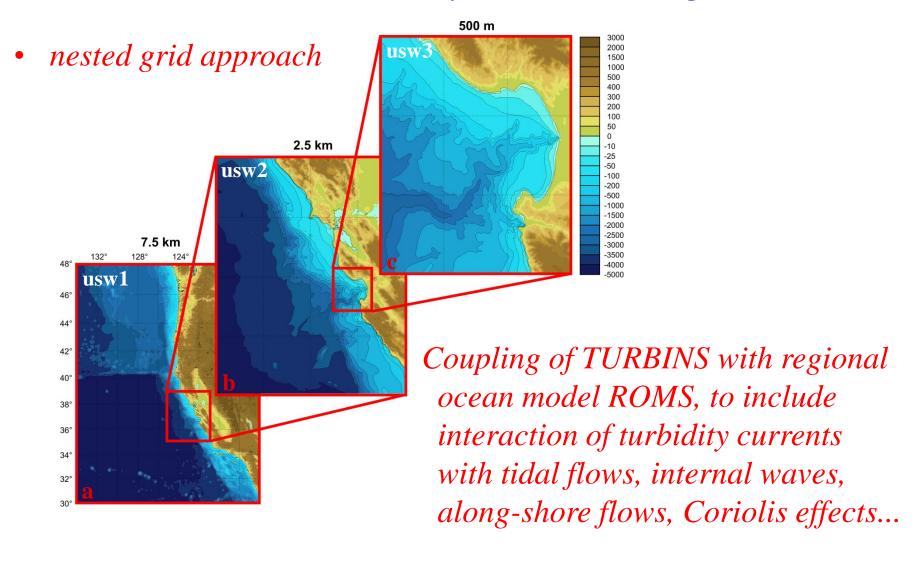
- While the large scale flow features are unique for every flow, the smallest scale flow features are similar for all turbulent flows → we may not have to resolve them, but instead may be able to model their main effect (energy extraction from large scales) by means of a turbulence model
- Two different approaches:
 - temporal averaging of governing equations →
 Reynolds-averaged Navier-Stokes (RANS) simulations
 - spatial averaging of governing equations → Large-eddy simulations (LES)

DNS simulation at $Re=10^3$ vs. LES simulation at $Re=2x10^5$ (with S. Radhakrishnan)



- higher Re current propagates faster, has more fine-scale structure
- similar flow structure, but large difference in bottom shear stress

Upscaling: Embedding high-resolution simulation within coarser resolution one (w. J. Syvitski, H. Arango, C. Harris)



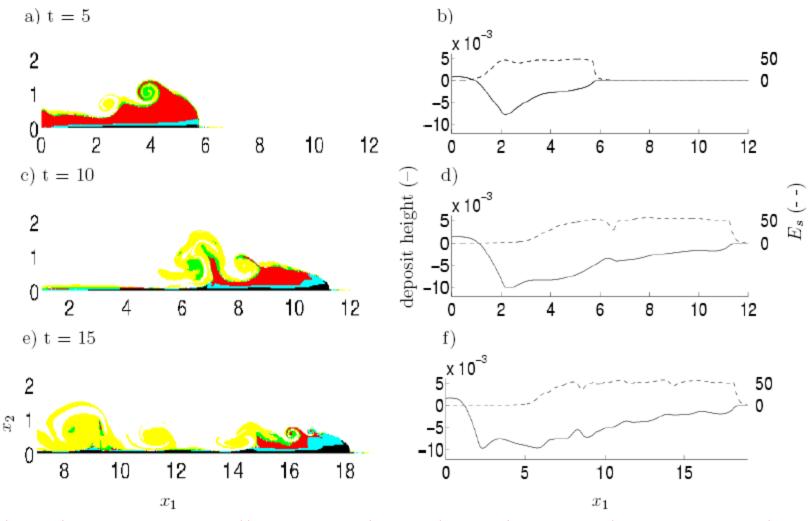
Erosion, resuspension of particle bed (with F. Blanchette, M. Strauss, B. Kneller, M. Glinsky)

Experimentally determined correlation by Garcia & Parker (1993) evaluates resuspension flux at the particle bed surface as function of:

- bottom wall shear stress
- settling velocity
- particle Reynolds number

Here we model this resuspension as diffusive flux from the particle bed surface into the flow

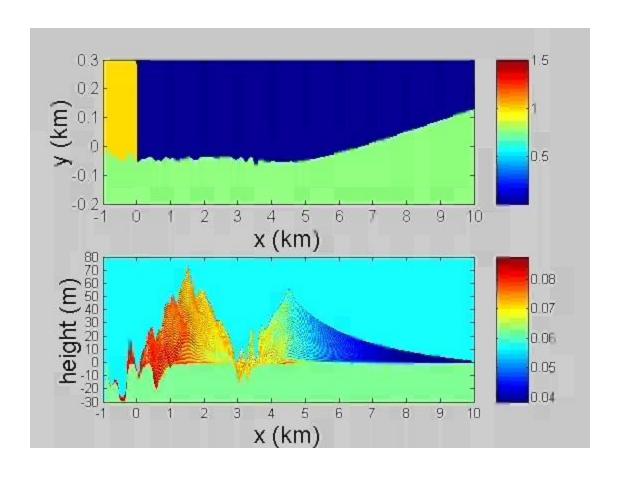
Erosion, resuspension of particle bed (cont'd)



• based on experimentally measured correlation between shear stress at the surface of the bed and an effective resuspension flux

Erosion, resuspension of particle bed (cont'd)

- multiple, polydisperse flows
- feedback of deposit on subsequent flows
- formation of ripples, dunes etc.



Erosion of sediment bed (Z. Borden, Y. Kanarska, M. Glinsky, E. Biegert)

- erosion models to date are mainly empirical, e.g. Garcia and Parker (1993), limited validity, not based on first principles → research at the microscopic level is needed to develop improved erosion models
- perform many-particle simulations, with the flow around each particle resolved
- employ model flows (Couette), subject sediment bed to increasing shear stress until erosion occurs
- study mechanics of erosion from first principles
- derive scaling laws for improved macroscopic, continuum erosion models

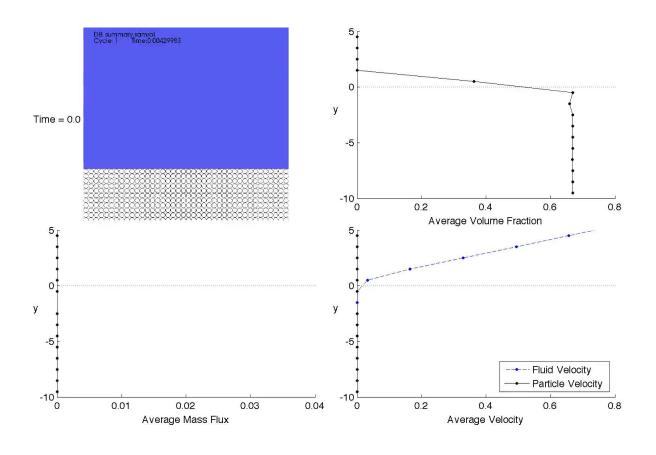
Erosion of particle bed: Couette flow (Z. Borden, L. Maurin)

2D simulation, Shields number = 0.16:

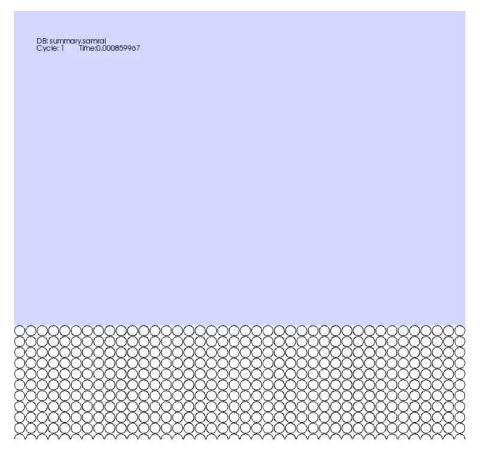


Borden, Maurin and Meiburg (2012)

• Extracting continuum information:

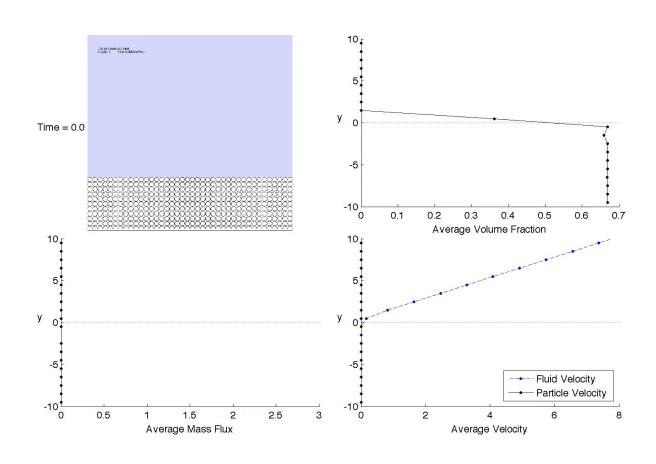


2D simulation, Shields number = 0.80:

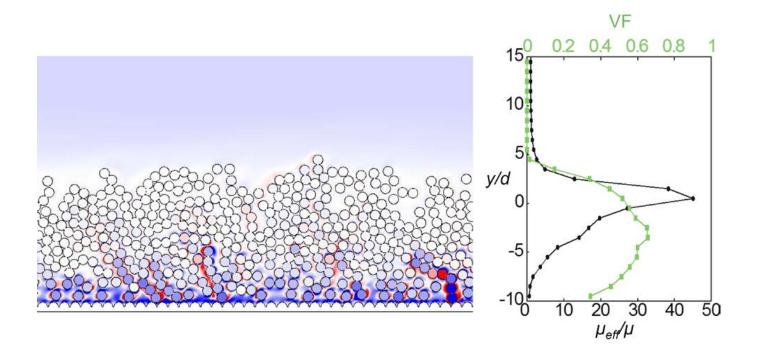


Borden, Maurin and Meiburg (2012)

Towards effective continuum boundary conditions:



Towards effective rheology and continuum boundary conditions:



• effective viscosity can increase by a factor of 50!



Particles of different sizes settling, interacting via collisions:

• study segregation mechanisms, spatial properties of resulting sediment bed

Grain-scale simulations

Advantages:

- accurately capture dynamics of individual grains
- provide very detailed information on grain/grain interactions
- can clarify mechanisms governing size segregation
- potential to extract effective rheology, and to upscale
- potential to analyze erosion of compacted vs. non-compacted sand
- potential to study the coupling between flow above and inside bed

Current challenges and questions:

- computationally very expensive
- limited to small scales, O(1,000) particles
- how relevant are the dynamics at the grain scale, compared to erosion of large chunks of sediment by large-scale energetic eddies?

Summary

- simulation tools for laboratory—scale, dilute depositional currents carrying small particles have contributed to our understanding
- extension to field scale via LES/RANS models is underway
- beginning to understand the physics behind bedform formation: ripples, dunes, antidunes, sediment waves, levees ...
- challenges: erosion, bedload transport, particle/particle collisions, hindered settling, concentration-dependent rheology
- alternative: two-fluid models, require empirical closure assumptions
- grain-scale simulations are beginning to contribute to our understanding of microscale phenomena
- limited to O(1,000) particles, but may provide information that will allow for progress with regard to upscaling
- importance of grain-scale phenomena for large-scale dynamics?

Outlook

- close gap between grain-scale and lab-scale ('mesoscale') \rightarrow upscaling
- extend lab-scale modeling to field scales via LES/RANS models
- need better understanding of current/bed interaction, including erosion, bedload transport, coupling between flow above the bed and inside the bed ...
- need better understanding of the influence of higher concentrations of particles: collisions, hindered settling, rheology
- progress will require coordinated advances in modeling (grain-scale, mesoscale, laboratory scale, field scale), laboratory measurements and field-scale observations
- it will be useful to define specific test cases to be analyzed from different perspectives

Acknowledgments

• B. Kneller, B. Hall, F. Blanchette, M. Glinsky, M. Strauss, M. Nasr-Azadani, S. Radhakrishnan, Z. Borden, E. Biegert