

Angular distribution of energy spectrum in two-dimensional β -plane turbulence in the long-wave limit

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Abstract

The time-evolution of two-dimensional decaying turbulence governed by the long-wave limit, in which $L_D/L \rightarrow 0$, of the quasi-geostrophic equation is investigated numerically. Here, L_D is the Rossby radius of deformation, and L is the characteristic length scale of the flow. As the degree of nonlinearity decreases, more energy accumulates in a wedge-shaped region where $||l| > \sqrt{3}|k|$ in the two-dimensional wavenumber space. Here, k and l are the longitudinal and latitudinal wavenumbers, respectively. When the degree of nonlinearity is decreased further, energy concentrates on the lines of $l = \pm\sqrt{3}k$. These results are interpreted based on the conservation of zonostrophy, which is an extra invariant other than energy and enstrophy and was determined in a previous study. Considerations concerning the appropriate form of zonostrophy for the long-wave limit and a discussion of the possible relevance to Rossby waves in the ocean are also presented.

1. Two-dimensional turbulence on a β -plane

- governed by Quasi-geostrophic equation.

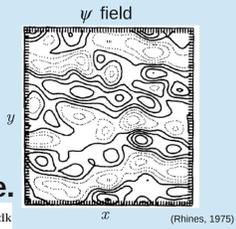
$$\frac{\partial}{\partial t} \left(\nabla^2 \psi - \frac{1}{L_D^2} \psi \right) + \beta \frac{\partial \psi}{\partial x} + \frac{\partial \psi}{\partial x} \frac{\partial \nabla^2 \psi}{\partial y} - \frac{\partial \psi}{\partial y} \frac{\partial \nabla^2 \psi}{\partial x} = 0$$

ψ : stream function L_D : deformation radius
 β : y-derivative of Coriolis parameter

- tends to have a zonally elongated structure.

- Effect of β -term

- prevents upscale energy cascade.
 - makes energy cascade anisotropic (Rhines effect).
 - becomes prominent when $L \sim L_D$.
- L : length scale of the flow (U : velocity scale of the flow)
 L_β : Rhines scale $L_\beta = \sqrt{U/\beta}$



- Zonostrophy

$$Z = \int f_k \varepsilon_k dk \quad \left\{ \begin{array}{l} \varepsilon_k = \frac{1}{2} (|k|^2 + L_D^{-2}) |\hat{\psi}_k|^2 \quad (\text{2D energy spectrum}) \\ f_k = \frac{|k|^2 + L_D^{-2}}{\beta k} \left(\tan^{-1} \left[\frac{l - k\sqrt{3}}{L_D |k|^2} \right] - \tan^{-1} \left[\frac{l + k\sqrt{3}}{L_D |k|^2} \right] \right) \end{array} \right.$$

- quasi-invariant for QG equation (Balk, 1991).
- used to explain anisotropic energy cascade (Balk, 2005).
- conserved when the nonlinearity is sufficiently weak.
- named by Nazarenko and Quinn (2009).

2. Purpose of this study

- Various parameter regimes

- There are six combinations of the three length scales L , L_D and L_β (Table 1)
- Regimes (5) and (6) ($L > L_D$, L_β) are not well-studied.

- We consider the long-wave limit ($L_D/L \rightarrow 0$) as an extreme case and:

- derive the asymptotic form of zonostrophy.
- investigate the anisotropy of energy cascade and check the conservation of zonostrophy (by numerical experiments).

No.	Regime	Anisotropy	MEMO
(1)	$L < L_D < L_\beta$		Move to regime (2) due to upscale energy cascade
(2)	$L_\beta < L < L_D$	YES	Predominance of zonal flow
(3)	$L < L_D < L_\beta$	NO	Checked by Okuno and Masuda (2003)
(4)	$L_D < L < L_\beta$	NO	Checked by Okuno and Masuda (2003)
(5)	$L_\beta < L_D < L$		Target of this study
(6)	$L_D < L_\beta < L$		Target of this study

Table 1

3. Long-wave limit

- Long-wave limit of quasi-geostrophic equation

Consider the quasi-geostrophic equation:

$$\frac{\partial}{\partial t} \left(\nabla^2 - \frac{1}{L_D^2} \right) \psi + \beta \frac{\partial \psi}{\partial x} + \frac{\partial \psi}{\partial x} \frac{\partial \nabla^2 \psi}{\partial y} - \frac{\partial \psi}{\partial y} \frac{\partial \nabla^2 \psi}{\partial x} = 0$$

In the long-wave limit ($L_D/L \rightarrow 0$), the operator in the first term expands as follows:

$$\begin{aligned} (\nabla^2 - L_D^{-2}) &= -L_D^{-2} (1 - L_D^2 \nabla^2) \\ &= -L_D^{-2} (1 + (L_D^2 \nabla^2)^2 + \dots)^{-1} \end{aligned}$$

After a proper Galilean transform and nondimensionalization, we get:

$$\frac{\partial \psi_*}{\partial t_*} + \gamma \frac{\partial \nabla_*^2 \psi_*}{\partial x_*} + \frac{\partial \psi_*}{\partial x_*} \frac{\partial \nabla_*^2 \psi_*}{\partial y_*} - \frac{\partial \psi_*}{\partial y_*} \frac{\partial \nabla_*^2 \psi_*}{\partial x_*} = 0$$

: Long-wave limit of QG equation

$$\begin{aligned} x &= Lx_*, & y &= Ly_*, \\ \psi &= UL\psi_*, \\ t &= \frac{L^3}{L_D^3 U} t_*, & \nabla^2 &= L^{-2} \nabla_*^2 \end{aligned}$$

Nondimensionalization

Dimensionless number

$$\gamma = \frac{\beta L_D^2}{U}$$

- (square of) the ratio of L_D to L_β
- does not depend on L
- considered as constant in time (since U is constant in time)
- introduced by Okuno and Masuda (2003)

- Long-wave limit of zonostrophy

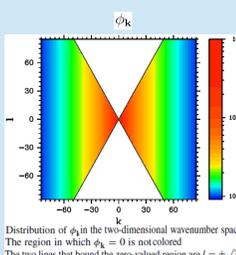
We found the following form:

$$Z = \iint \phi_k \varepsilon_k dk \quad \left\{ \begin{array}{l} \varepsilon_k = \frac{1}{2} |\hat{\psi}_k|^2 \quad (\text{2D energy spectrum}) \\ \phi_k = \begin{cases} 1/|k| & (|l| < \sqrt{3}|k|) \\ 0 & (|l| > \sqrt{3}|k| \text{ or } |k| = 0) \\ 1/(2|k|) & (|l| = \sqrt{3}|k| \text{ and } |k| \neq 0) \end{cases} \end{array} \right.$$

- conserved when the nonlinearity is sufficiently weak (i.e. when $\gamma \gg 1$)
- if conserved, energy should accumulate in a wedge-shaped region W in 2D wavenumber space:

$$W = \{(k, l) \mid \sqrt{3}|k| < |l|\}$$

(corresponding to the blank region in upper-right figure.)



Distribution of ϕ_k in the two-dimensional wavenumber space. The region in which $\phi_k = 0$ is not colored. The two lines that bound the zero-valued region are $l = \pm\sqrt{3}k$.

4. Numerical experiments

- Experimental setup

- Governing equation:

- Long-wave limit of QG eq. + hyper viscosity

$$\frac{\partial \psi}{\partial t} + \gamma \frac{\partial \nabla^2 \psi}{\partial x} + \frac{\partial \psi}{\partial x} \frac{\partial \nabla^2 \psi}{\partial y} - \frac{\partial \psi}{\partial y} \frac{\partial \nabla^2 \psi}{\partial x} = (-1)^p \nu_p \nabla^{2p} \psi$$

(Freely decaying turbulence) $\nu_p = 10$, $\nu_p = 5 \times 10^{-35}$

Domain	$2\pi \times 2\pi$, double periodic
Discretization	Space: Fourier spectral method ($K_r=85$) Time: 4th order Runge-Kutta ($\Delta t = 2 \times 10^{-7}$)
Parameter	$\gamma = 0, 0.25, 1, 5, 20$
Initial condition	random Ψ field

- Results

- 2D energy spectrum (average of 41 ensemble members) and their angular distribution

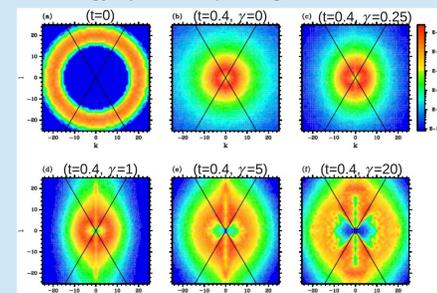


FIG. 3. (a) Two-dimensional energy spectrum at $t=0$ within a rectangular domain of $|k|, |l| \leq 25$ averaged over 41 ensemble members. (b)-(f) Same as (a) except that at $t=0.4$ for $\gamma = 0, 0.25, 1, 5$, and 20 , respectively. E→x indicates 10^{-4} . In each figure, the lines of $l = \pm\sqrt{3}k$ are drawn for reference.

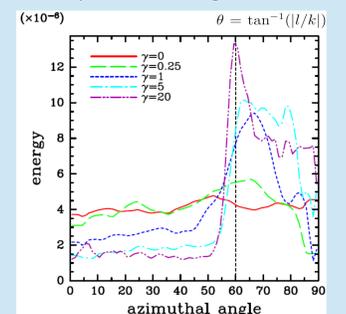


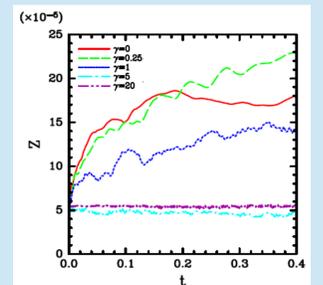
FIG. 4. Angular distributions of energy spectra corresponding to the final ($t=0.4$) states shown as Figs. 3(b)-3(f).

- Clear contrast across the boundary ($\theta = 60^\circ$)

- $\gamma=0$: Isotropic
- $\gamma=0.25$: Slightly elongated along the l -axis
- $\gamma=1, 5$: Accumulation in the wedge-shaped region W
- $\gamma=20$: Concentration around $\theta = 60^\circ$ ($\theta = \tan^{-1}(l/k)$) (Connaughton et al., 2011)

- Time-evolution of zonostrophy

- Zonostrophy is well-conserved for larger value of γ ($=5, 20$).



Time-evolutions of zonostrophy (Z) for five γ s.

5. Summary

- In the present study, we studied 2D turbulence governed by the long-wave limit of quasi-geostrophic equation and revealed:

- The asymptotic form of zonostrophy.

(by numerical experiments)

- The relationship between the parameter γ and anisotropy of 2D energy spectrum.

- $\gamma > 1$: Accumulation in the wedge-shaped region W .
- $\gamma \gg 1$: Concentration around $\theta = 60^\circ$.

- * Accumulation in wedge-shaped region W can be explained by conservation of zonostrophy.

6. Discussion

- 1st baroclinic Rossby waves in the mid-latitude ocean

- wave length: several hundreds – thousand (km)
- characteristic velocity scale: $U \sim 1$ (cm/s)
- observed in the mid-latitude ocean ($L_D \sim$ several tens (km))
- their wavenumber vector typically deviate from the zonal by $50^\circ - 80^\circ$ (Glazman and Weichman, 2005)

- For example, for 25° latitude, $\beta L_D^2 = 5.8$ (cm/s) and we get:

$$\gamma = \frac{\beta L_D^2}{U} \sim 5$$

which satisfies the condition for anisotropy ($\gamma > 1$).

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