# Submesoscale sea ice-ocean interactions in marginal ice zones



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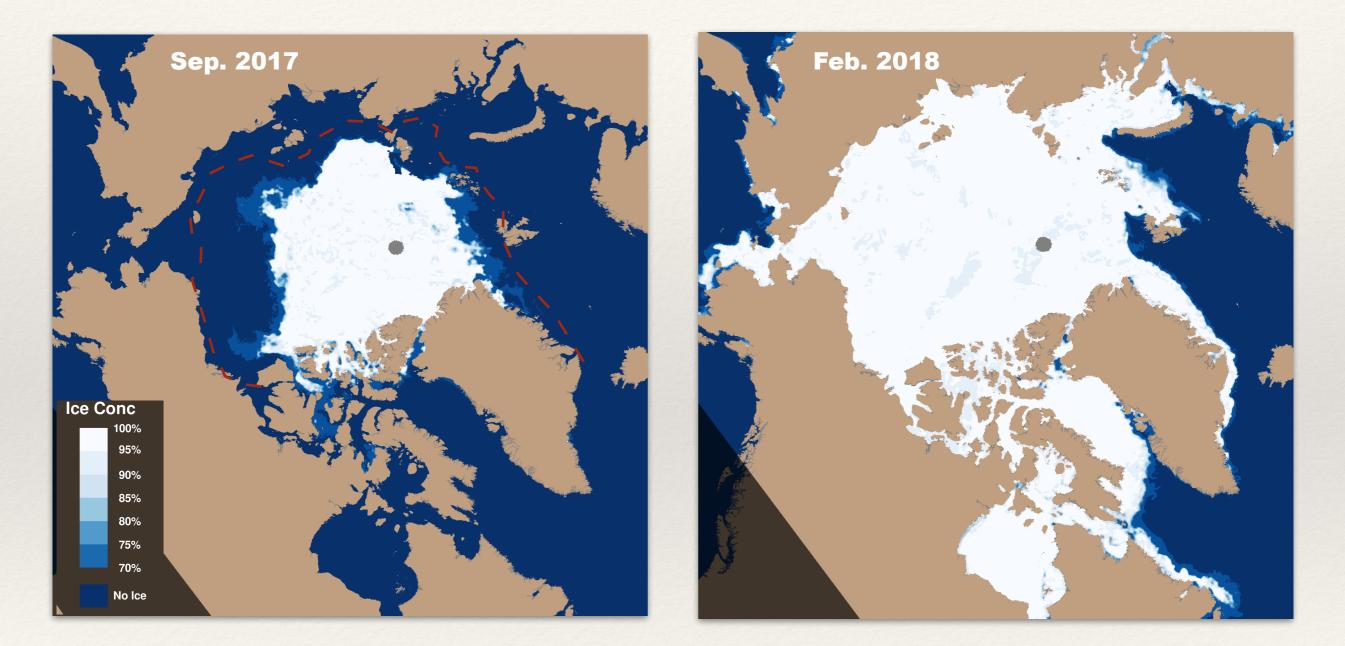
#### Collaborators: A. Thompson (Caltech), D. Menemenlis (JPL), R. Fajber (U. Toronto)

Sponsors: Stanback Fellowship and Davidow Discovery Fund (Caltech)



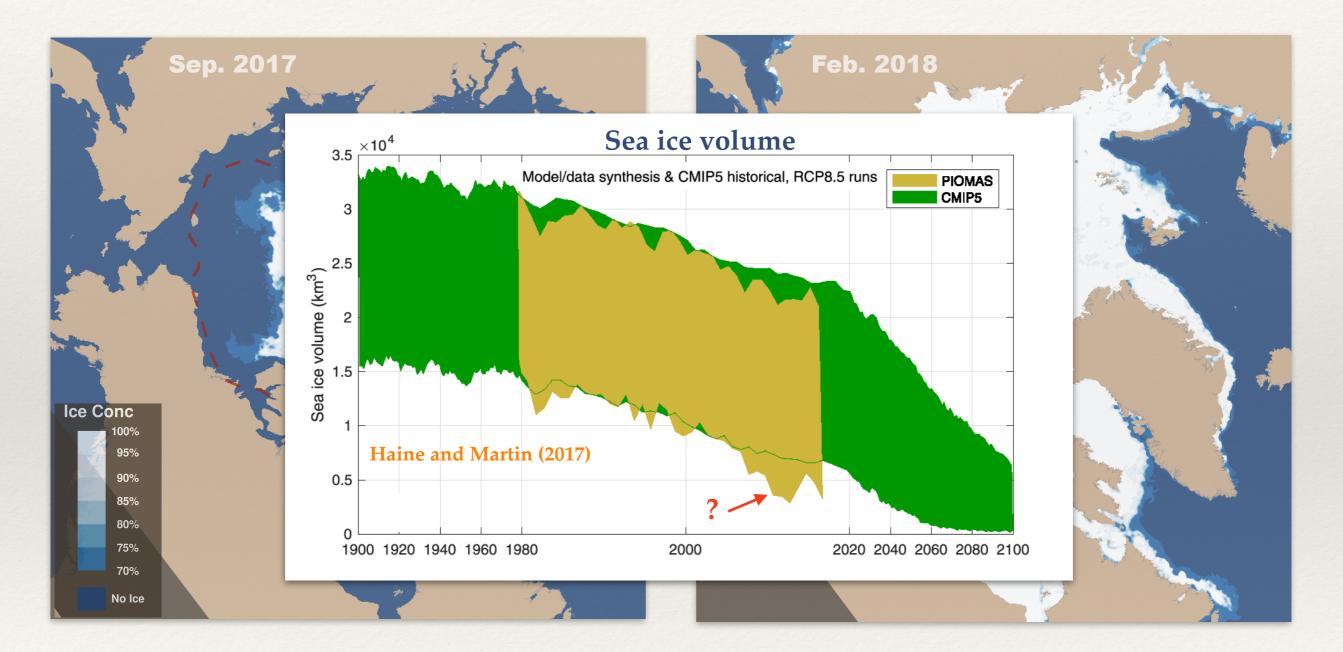
worldview.earthdata.nasa.gov

## Arctic sea ice state



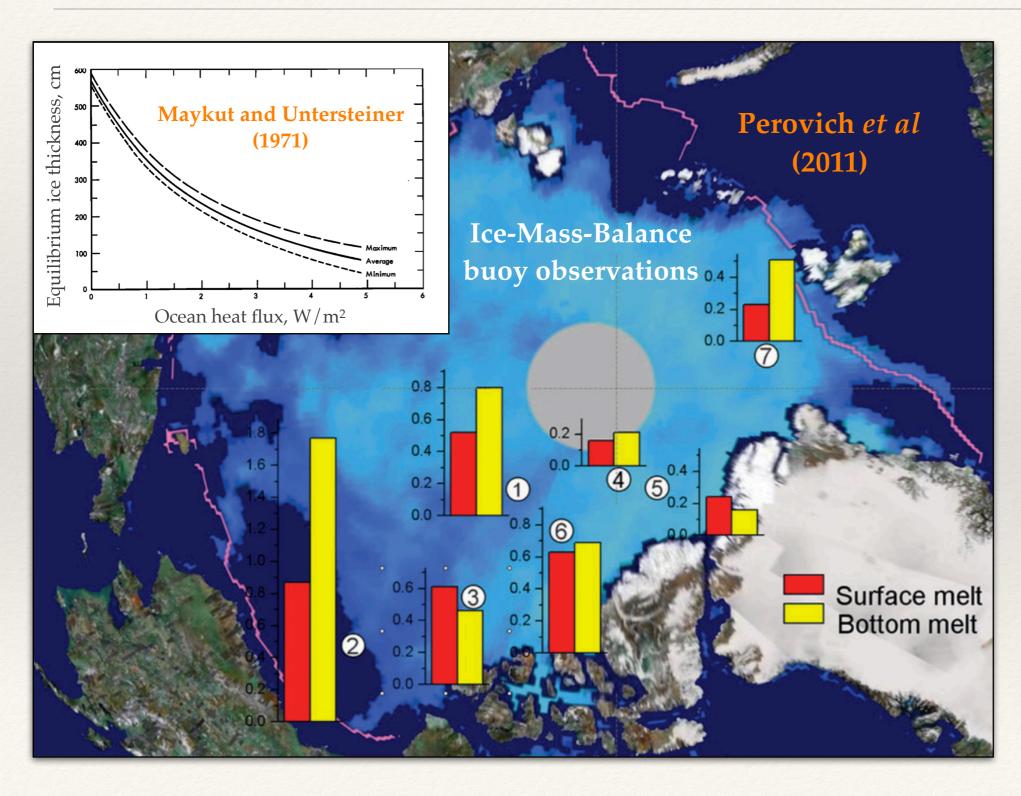
Data source: MASAM2, Daily 4-Km Arctic Sea Ice Concentration; National Snow and Ice Data Center, CIRES, University of Colorado, Boulder.

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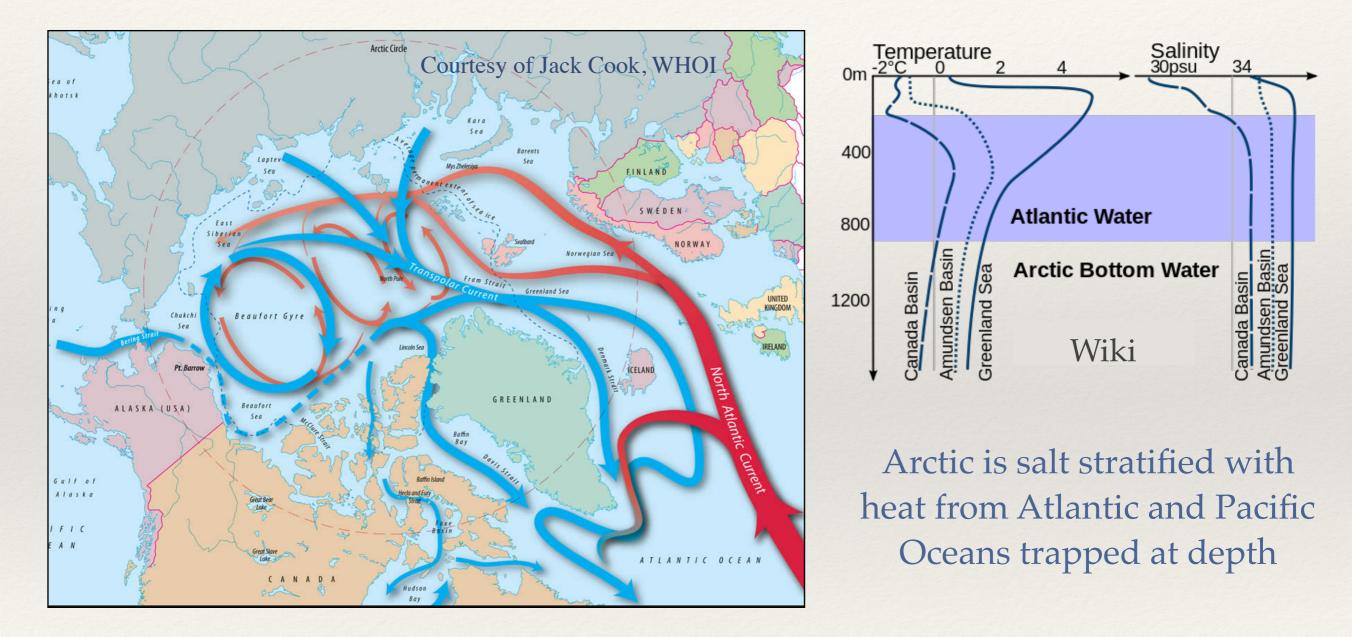
## Critical role of ocean heat in Arctic sea ice melt



Heat flux that can explain the sea ice loss during the past few decades is ≈1 W/m<sup>2</sup> (Kwok & Untersteiner, 2011)

Seasonal to interannual predictions have largest errors in MIZs (**Tietsche et al.**, **2014**) with extent predictability for only several months (**Stroeve et al.**, 2014).

## Where does the Arctic Ocean heat come from?



Heat stored in the halocline is an important source during winter

What is the role of Arctic eddies in heat advection?

LLC4320 Resolution < 1km

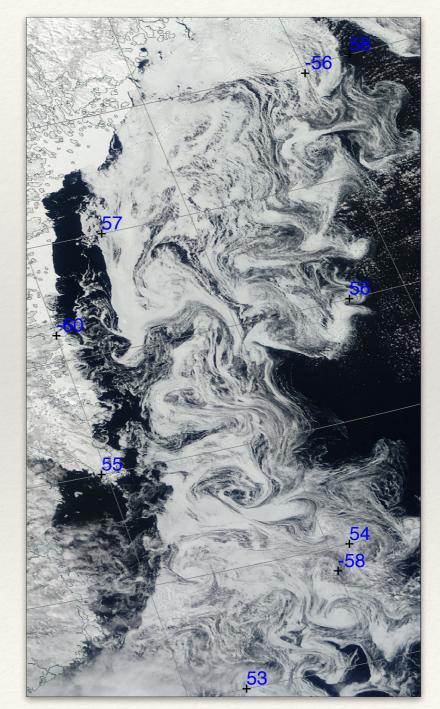
 $\zeta / f @ 80m$ 

### Heavy footprints of oceanic **fronts**, **eddies** and **filaments** on **sea ice** distribution in marginal ice zones

Beaufort Gyre from SAR imagery

#### Labrador Current (MODIS) Fram St. (MODIS)



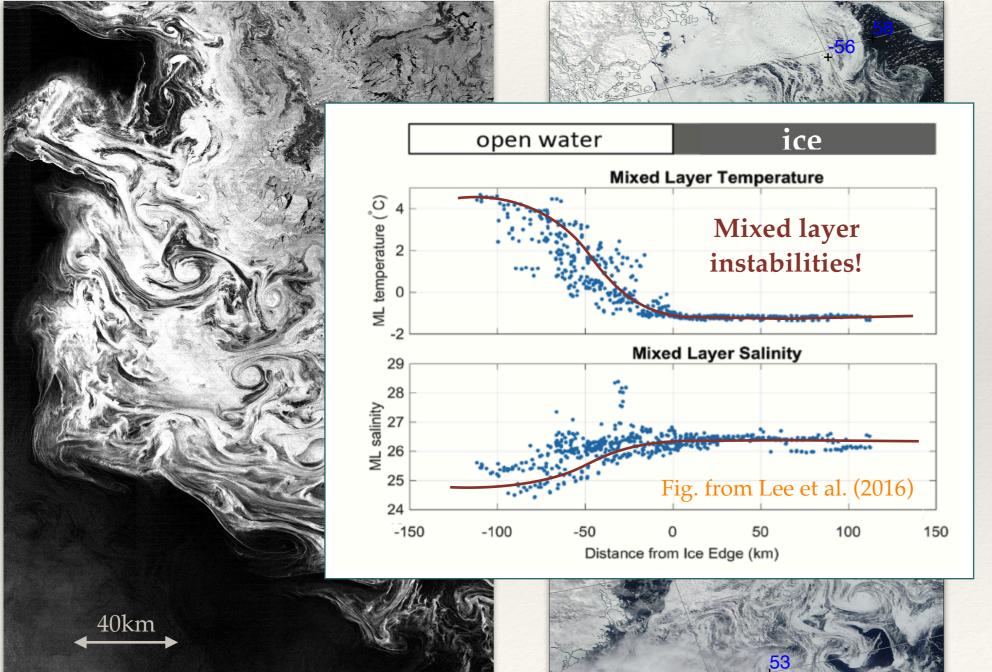




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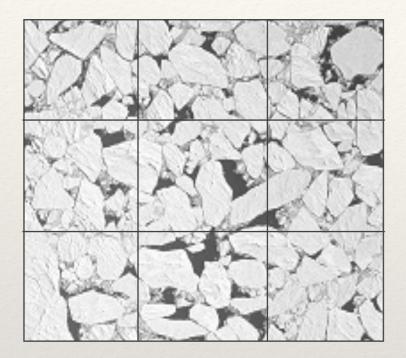


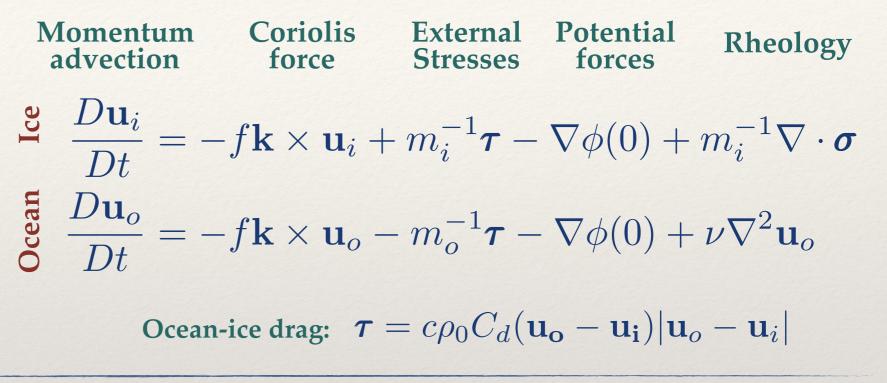


Part I. Sea ice-ocean interactions at eddy scales

Can eddies efficiently redistribute sea ice and enhance vertical heat fluxes? How robust is this mechanism?

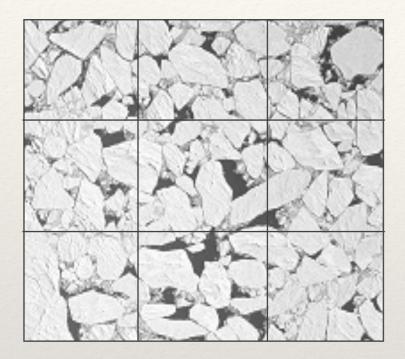
## Mechanical sea ice-ocean interactions

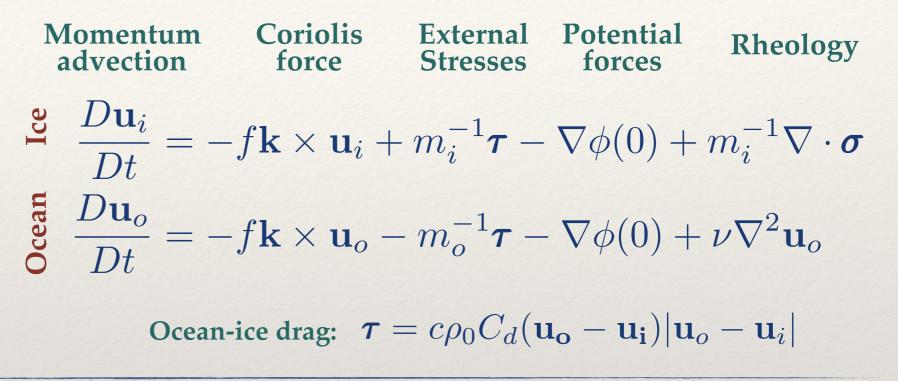




Sea ice rheology: e.g. visco-plastic shear-thinning fluid (Hibler III,1979)

## Mechanical sea ice-ocean interactions





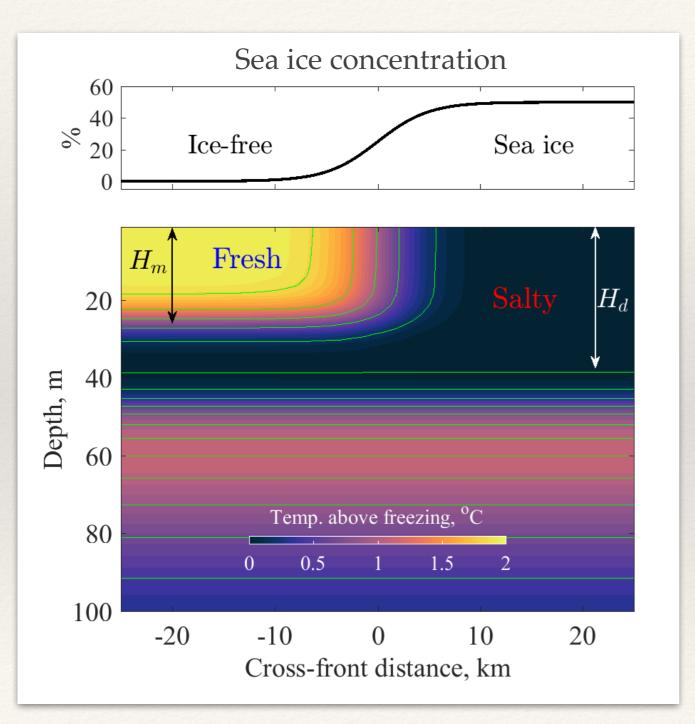
Sea ice rheology: e.g. visco-plastic shear-thinning fluid (Hibler III,1979)

$$\nabla \cdot \sigma \sim m_i f(u_i - u_o) \sim \tau \quad \Rightarrow \quad \frac{\zeta_i - \zeta_o}{\zeta_o} \sim \frac{m_i}{R_d \operatorname{Ro} C_d} \sim \frac{O(0.1)}{\operatorname{in MIZs}}$$

Manucharyan & Thompson (2017)

Sufficiently low-concentrated sea ice over submesoscale eddies must mimic the upper-ocean vorticity!

## Simulations of ice-covered mixed layer instabilities



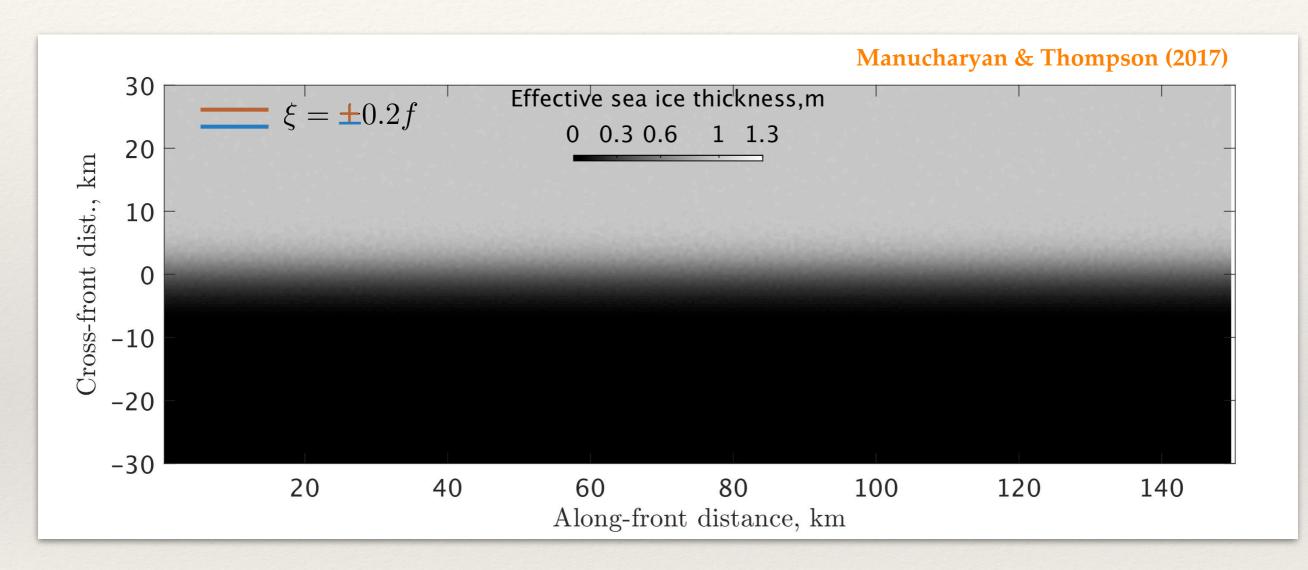
<u>Hypothesis</u>: instabilities of meltwater fronts in MIZs can strongly affect sea ice dynamics

• **MITGCM:** simulations of sea ice dynamics (as a visa-plastic fluid) over an eddying mixed layer.

Initial frontal conditions are idealizations of the observed fronts.

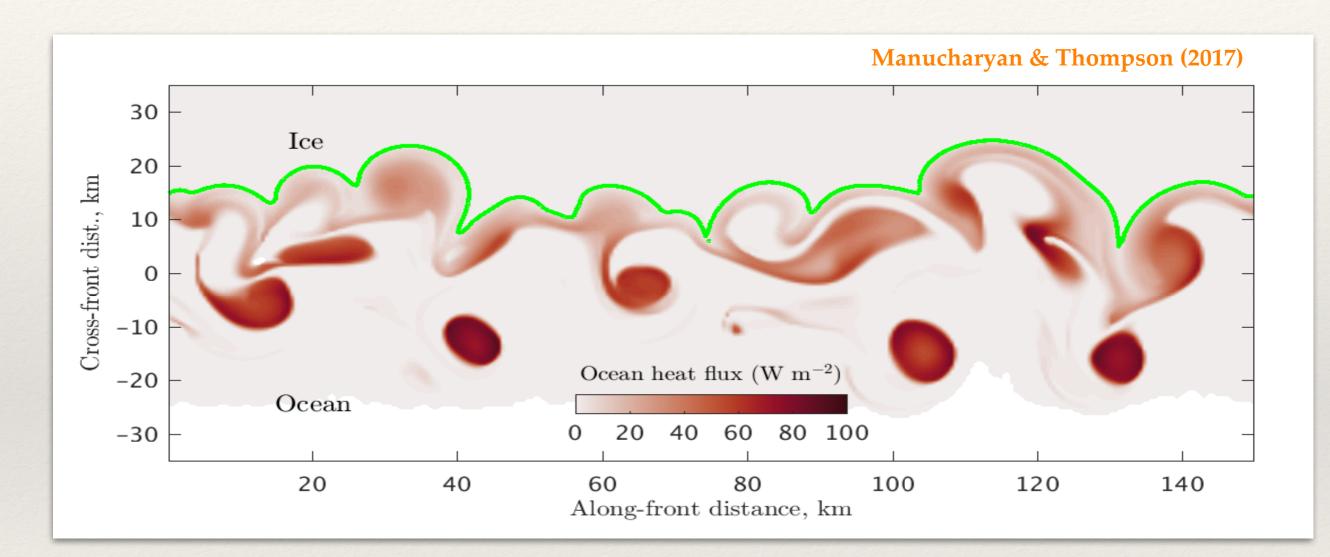
Grid of 0.5 km allows the development of frontal instabilities and ice-ocean stresses that redistribute sea ice

## Simulations of ice-covered mixed layer instabilities



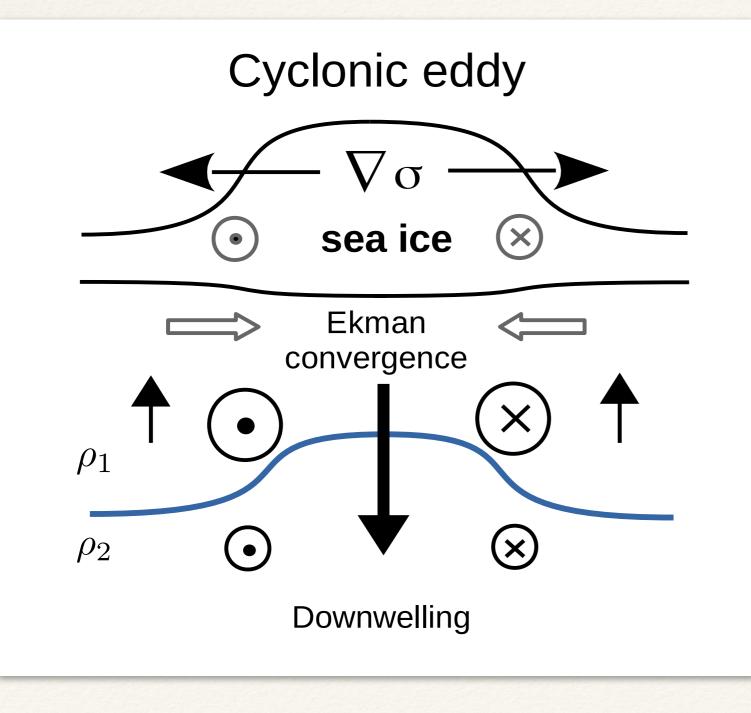
- Marginal ice zone expands due to an intense lateral sea ice transport. Sea ice is not a passive tracer.
- Cyclonic eddies (~10km, 0.2 m s<sup>-1</sup>) accumulate the sea ice while anticyclones repel it.
- Vertical velocities are O(10 m day<sup>-1</sup>) and ice-ocean heat fluxes reach O(100 W m<sup>-2</sup>)!

### Enhanced sea ice-ocean heat fluxes in marginal ice zones



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## Sea ice accumulation in cyclonic eddies and filaments



**Constitutive**  $P(c) = P_0 c^{\alpha}$ law:

> Mass conservation:  $\partial_t c + \partial_y (v_i c) = 0$

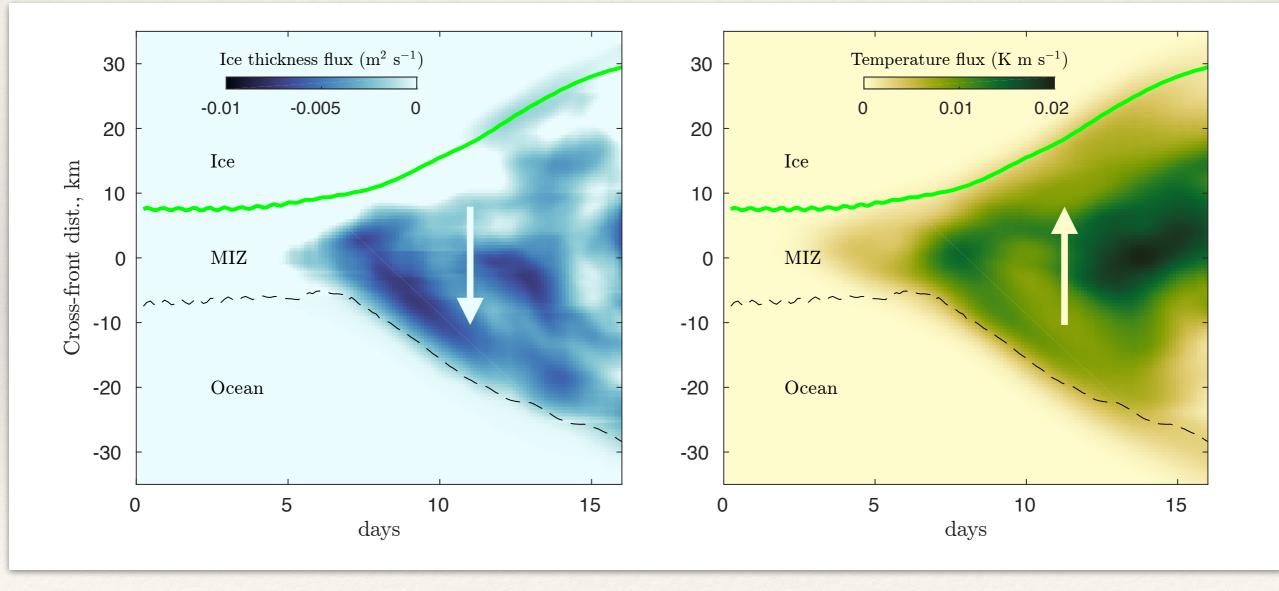
Force balance:  $\rho C_d c(v_i - v_o) + h P_0 \partial_y c^\alpha = 0$   $\downarrow$ Non-linear adv.+diffusion:  $\partial \hat{c}_t = -\partial_y [v_0(y)\hat{c} - \partial_y \hat{c}^\alpha]$ 

Fajber et al (in prep.)

## Lateral heat and sea ice transport across the MIZ

#### **Ice volume transport:** $< v'h'_{eff} >$

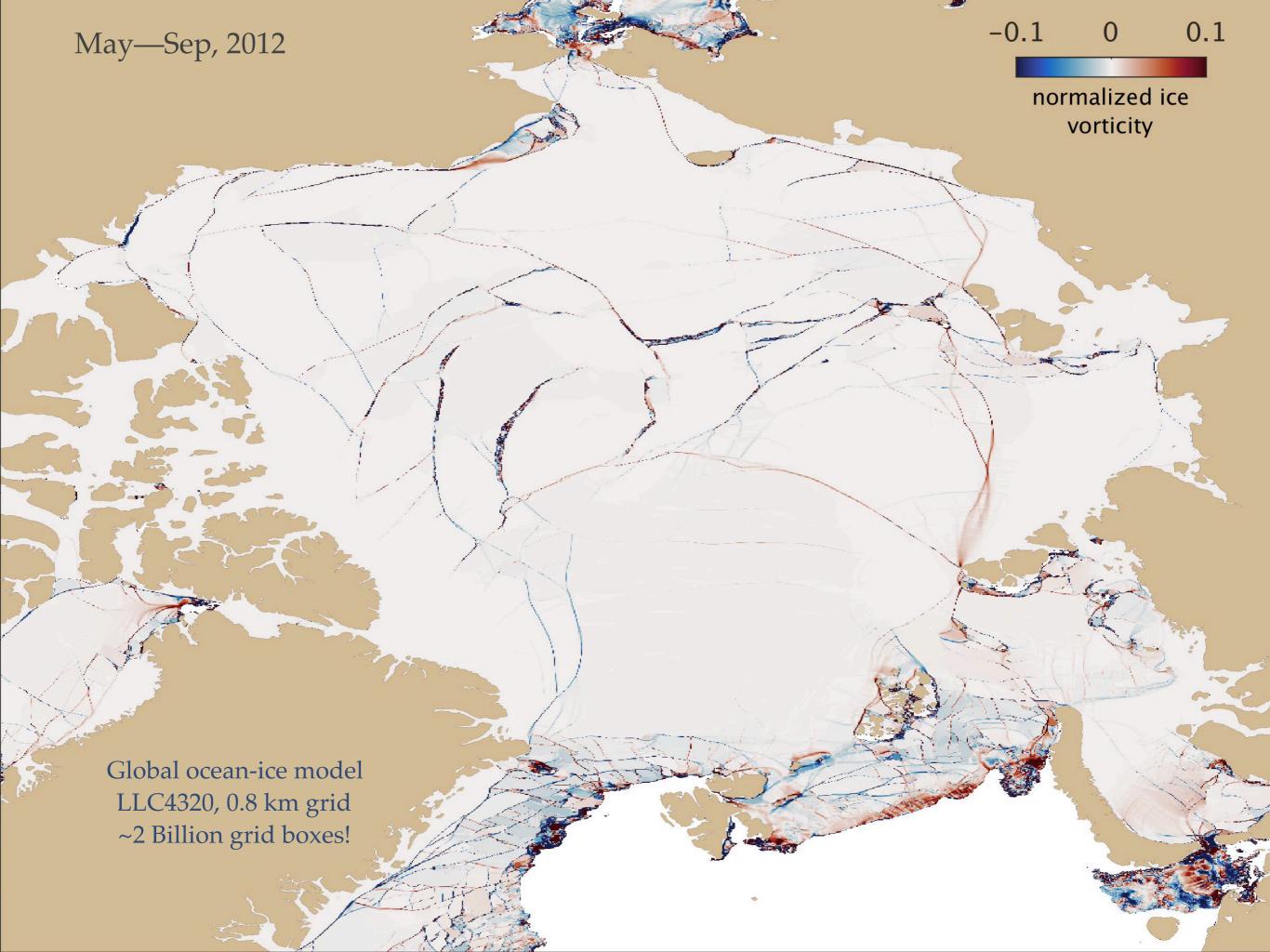
#### Surface heat transport: < v'T' >



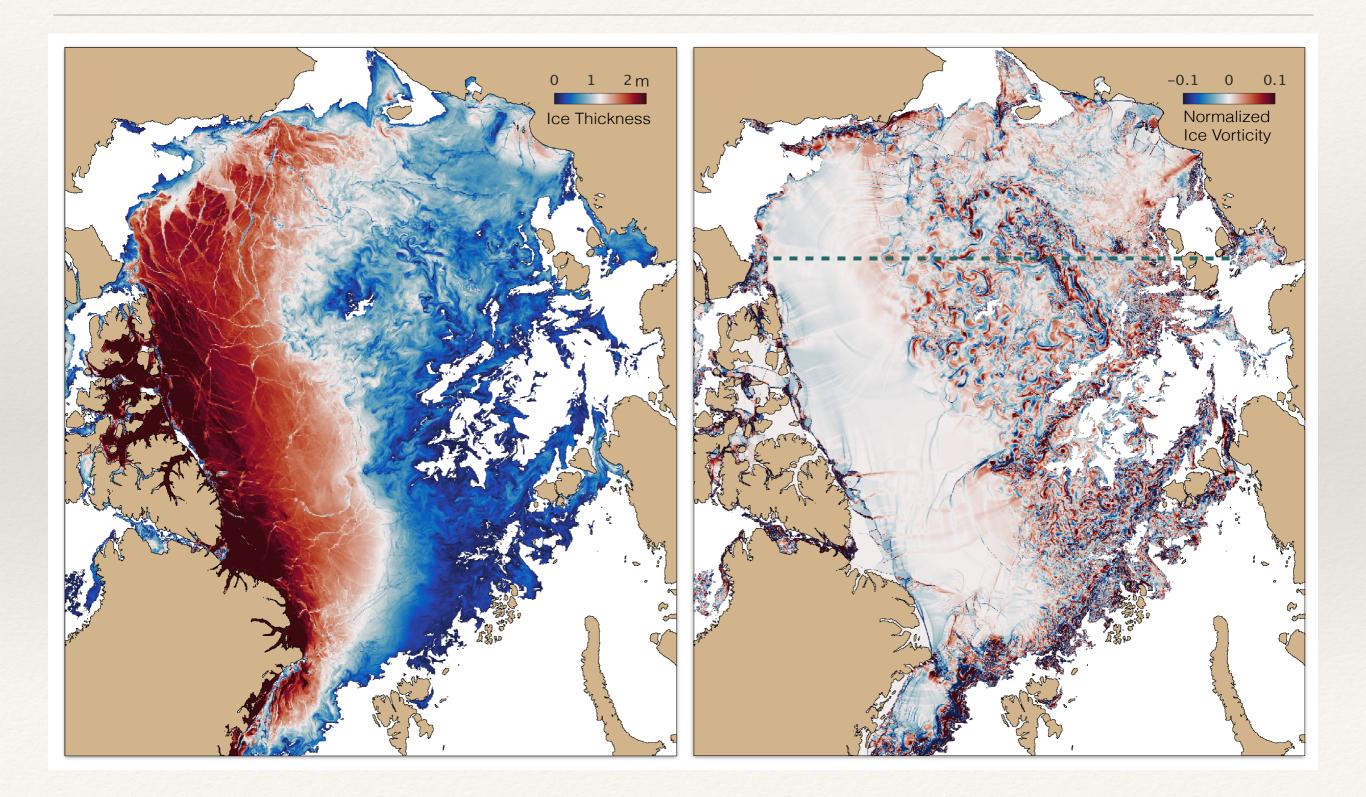
Tracer diffusivity  $K=O(200m^2/s)$ ; scales linearly with initial frontal Rossby number.

# **Part II.** Arctic-wide influence of mesoscale and submesoscale ocean variability on sea ice

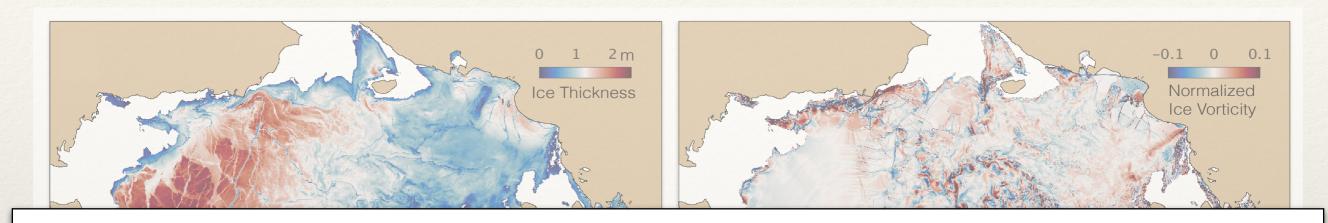
Is there a significant mesoscale/submesoscale variability in the Arctic? How does it affect the sea ice dynamics?

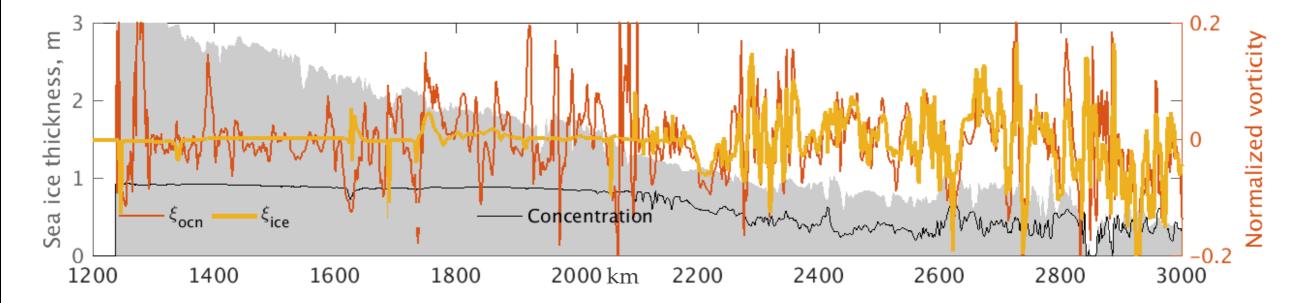


## Thickness {seaice} Vorticity

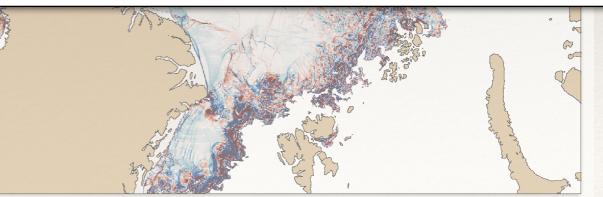


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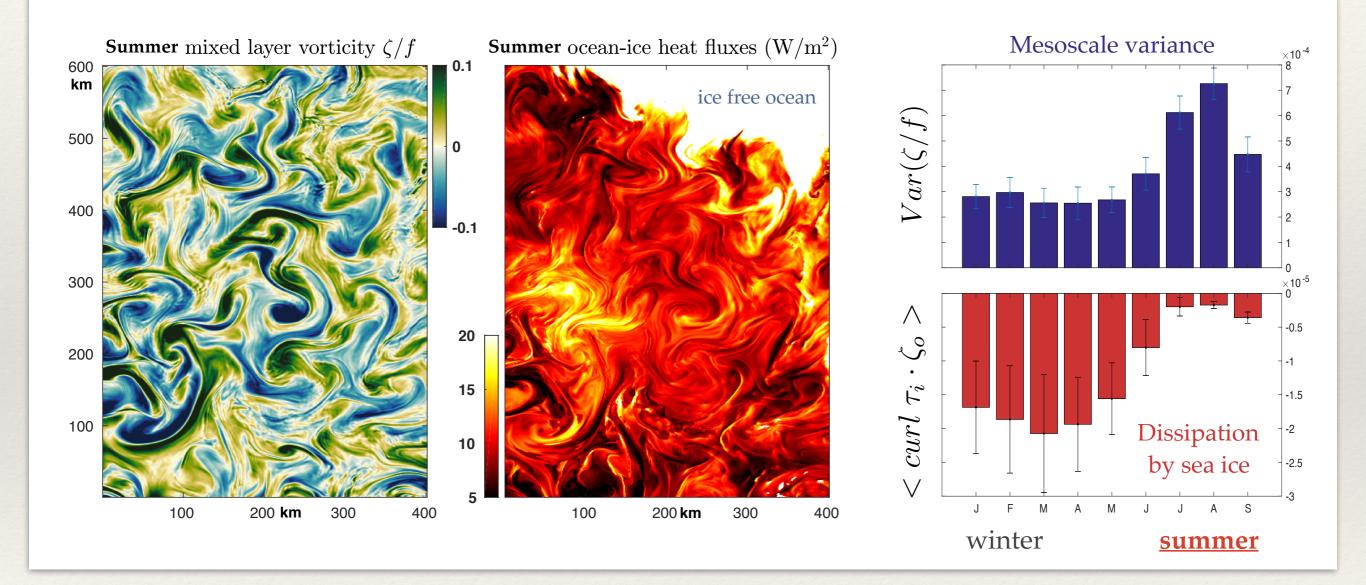






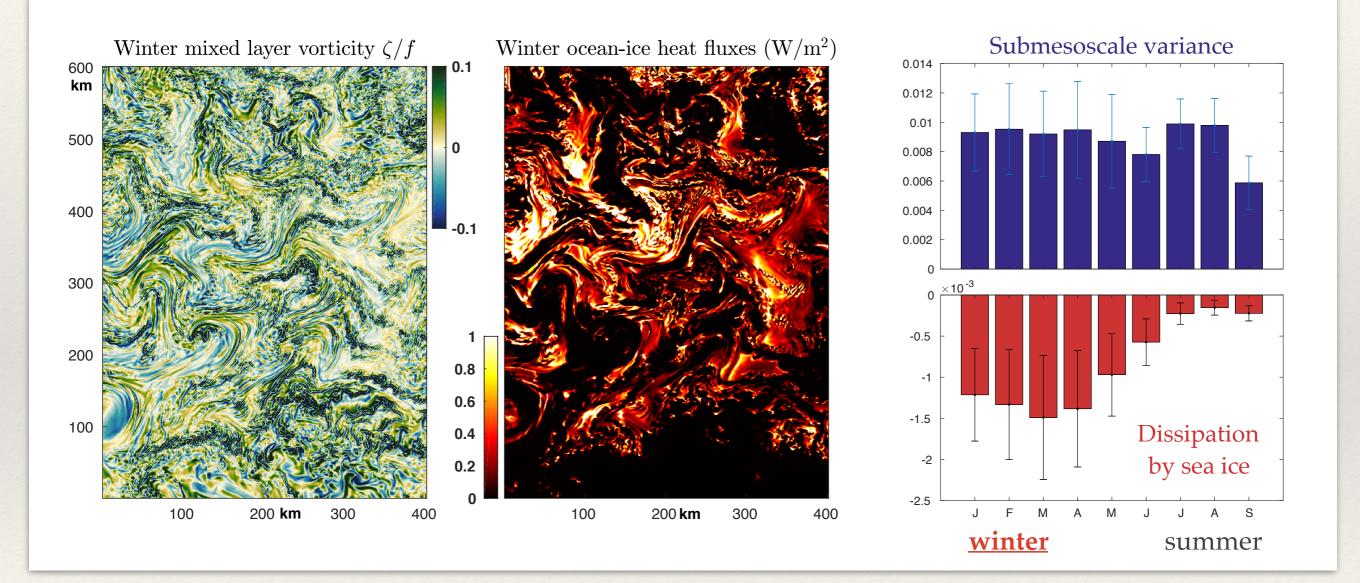


## Eddy-enhanced ocean-ice heat fluxes: summer



Surface expression of mesoscale eddies is strongest in the summer when dissipation by sea ice is minimized. Ocean-ice heat flux is a stirred field!

## Eddy-enhanced ocean-ice heat fluxes: winter



Despite low Ro~0.2, winter-time submesoscale activity constitutes about 80% of the average 2.4 W/m<sup>2</sup> ocean-ice heat fluxes Arctic-wide.

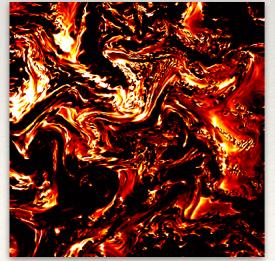
## Summary of sea ice-ocean interactions

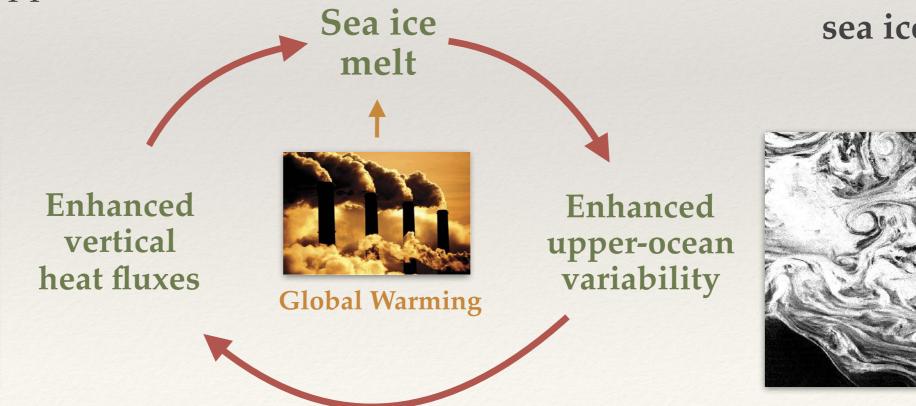
MLI is prominent in MIZs, strongest frontogenesis occurs in summer and convective instabilities in winter.

Sea ice gets trapped in cyclones and gains its vorticity, dramatically decreasing the upper-ocean dissipation.

Sea ice dynamics is getting more 'turbulent' because of weaker eddy dissipation in the upper ocean.

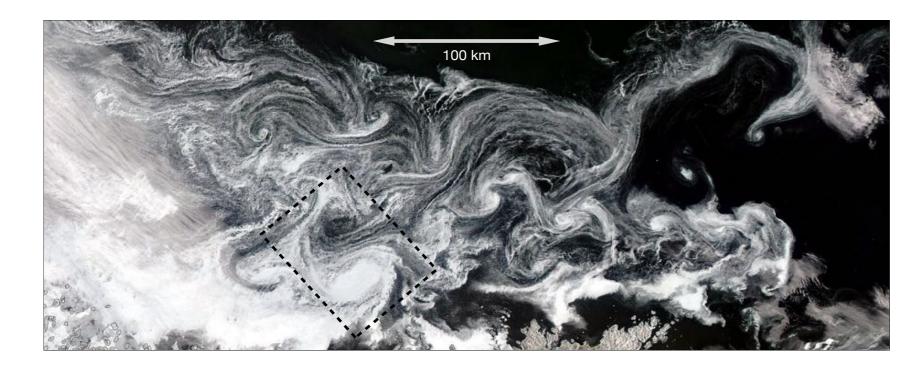
Increased eddy energy may contribute to the enhanced ocean heat fluxes and accelerated sea ice melt.





## More goods?

## Seeing the ocean through sea ice

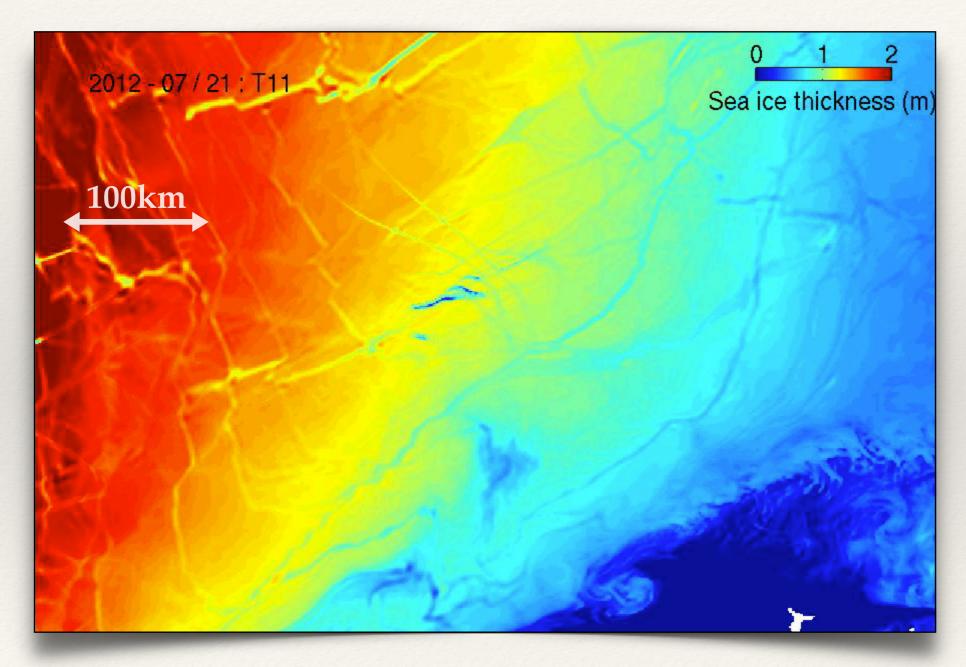


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$$u_o = F(u_i, \sigma) ?$$

- What is the relation between sea ice and ocean velocities?
- Under which atmospheric conditions the ocean vorticity field is reflected in sea ice concentrations?
- What can we learn about submesoscale ocean dynamics from sea ice observations?

## Addressing scale interactions in sea ice dynamics: eddies, internal waves, and mixing hotspots



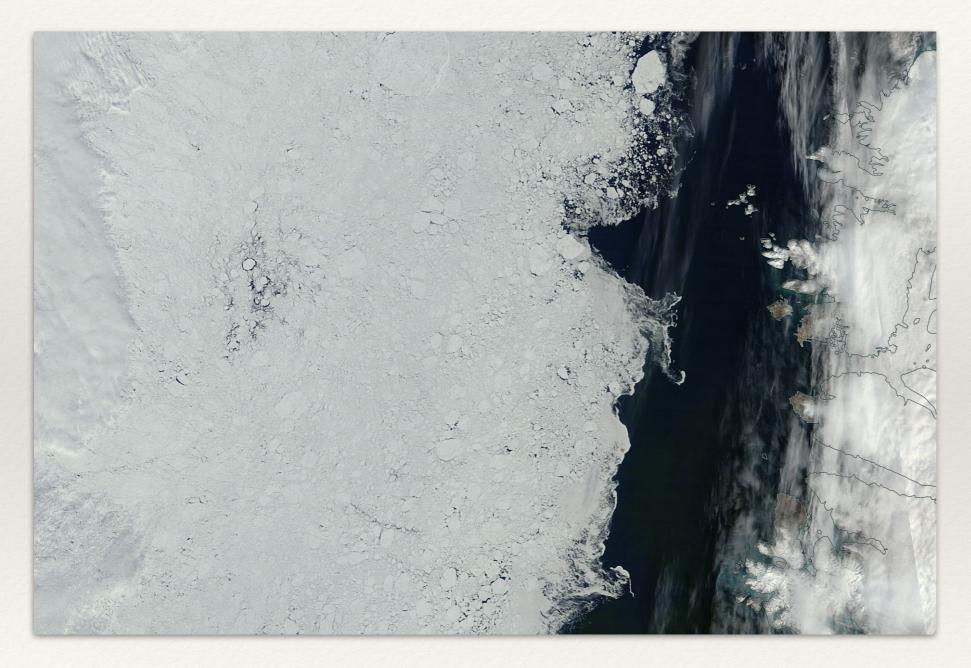
Internal waves break up the sea ice exposing it the open ocean to atmospheric fluxes

Mixing hotspots in Arctic Ocean affect halocline thickness and vertical heating

Eddies and waves become more energetic because of the changing damping

Sea ice dynamics North of Svalbard (MITGCM, ~1km res.)

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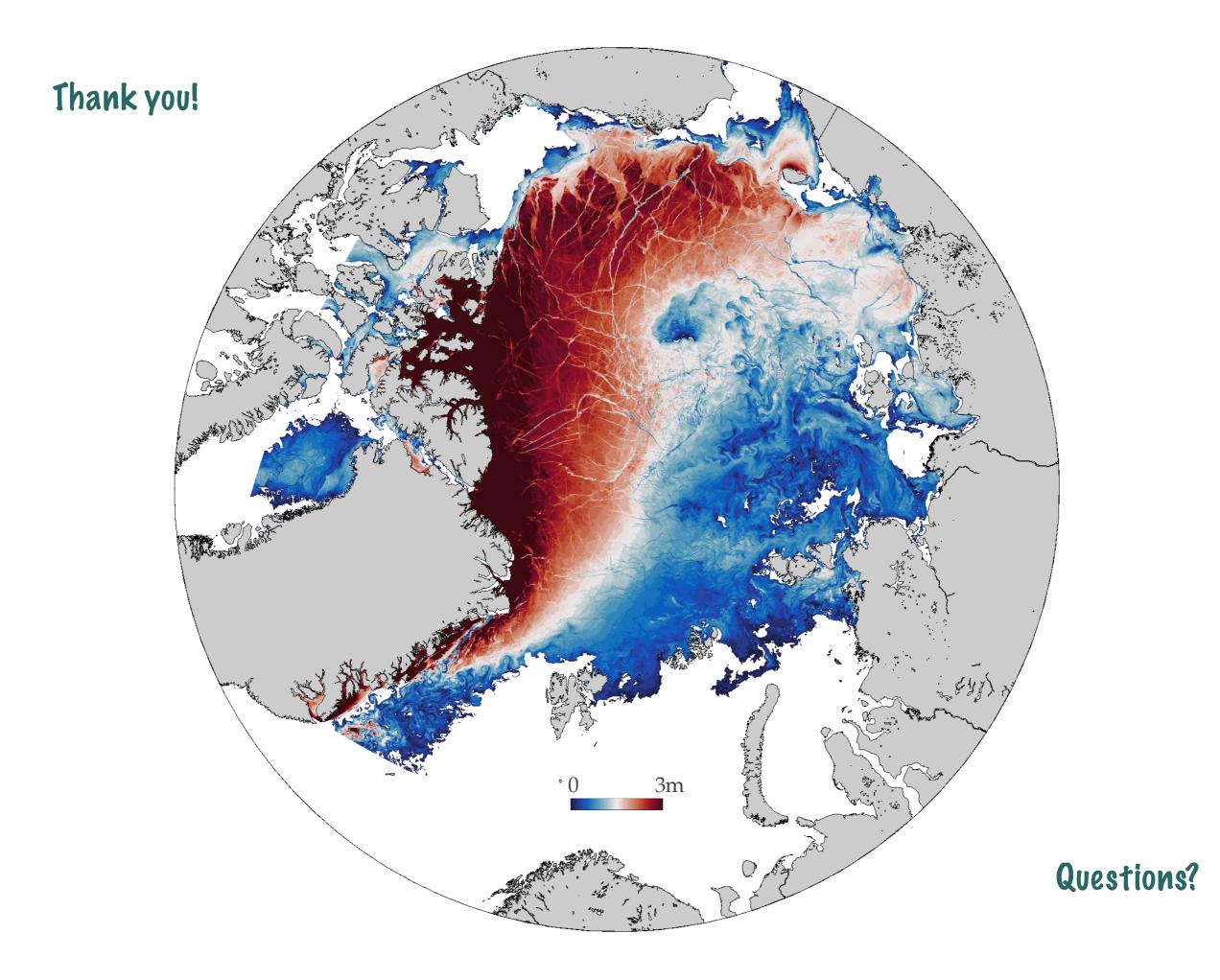


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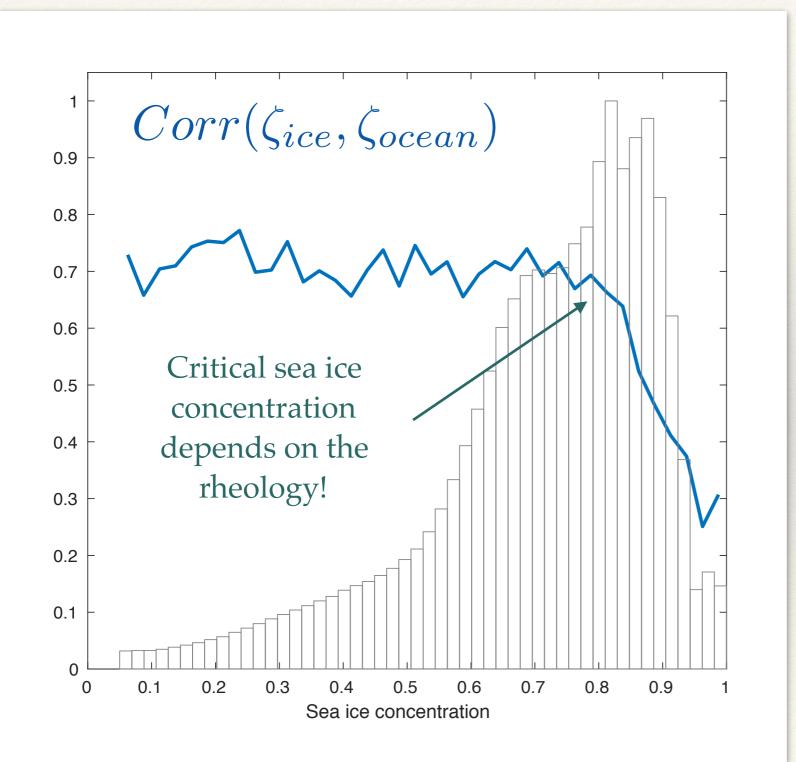
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## Critical influence of sea ice on upper-ocean vorticity

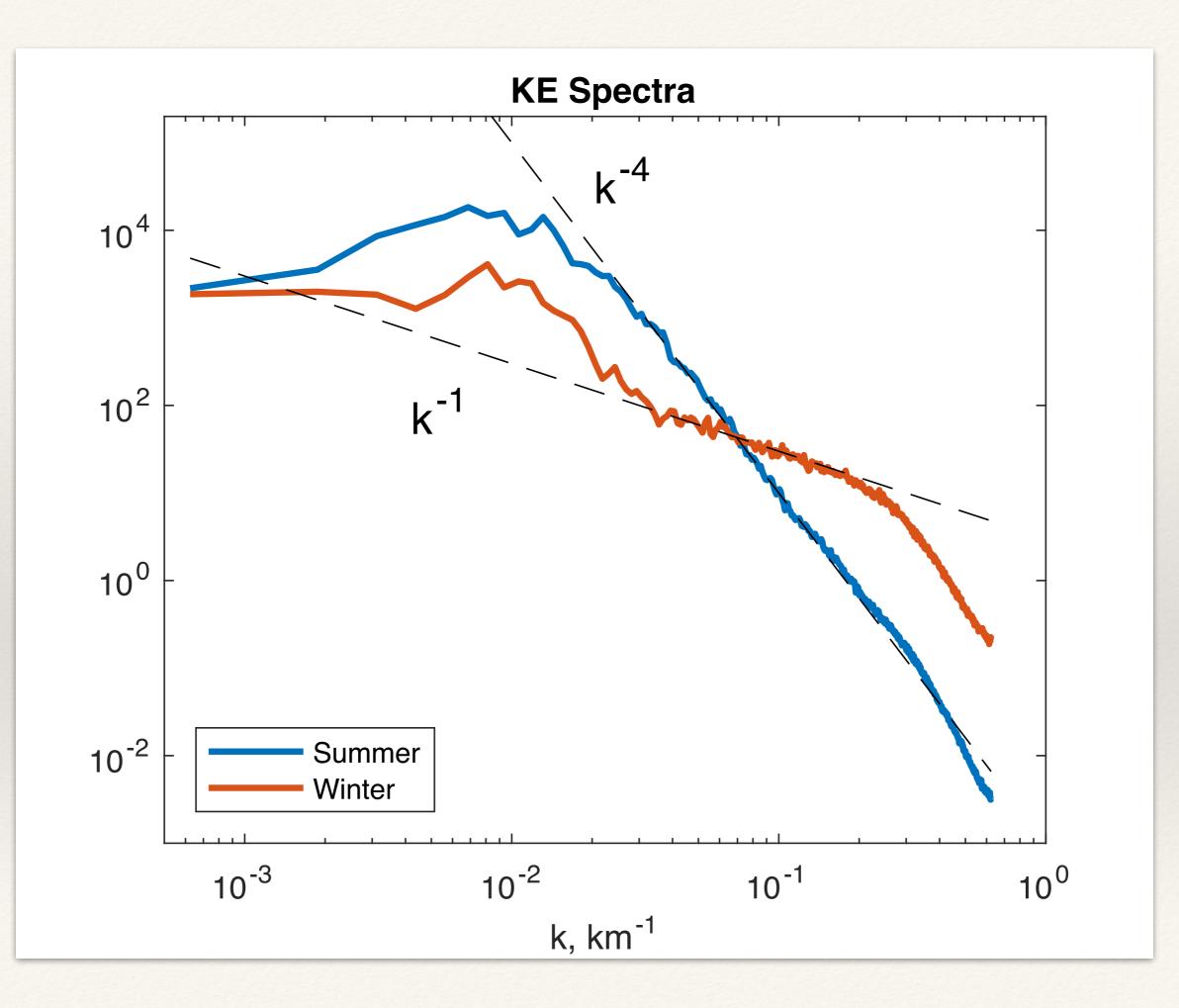


Ekman eddy spindown

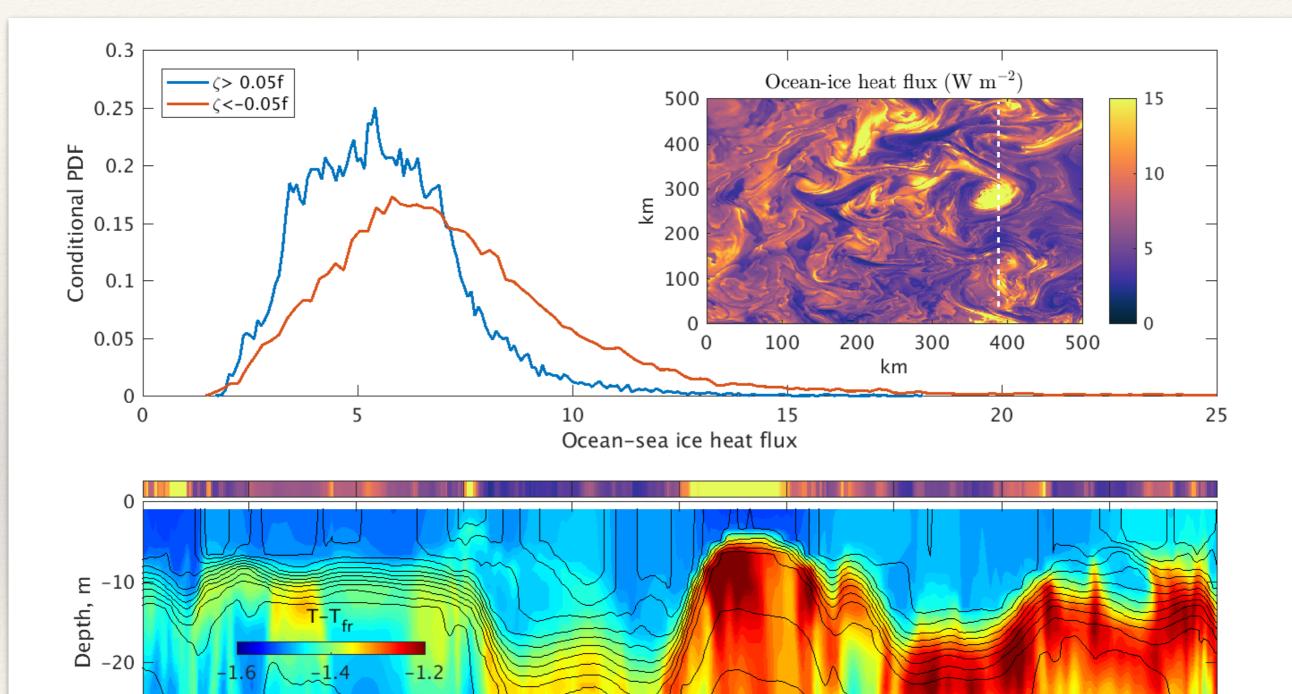
$$\nabla \times \tau_i \sim (\zeta_i - \zeta_o)$$

Vorticity dissipation by ice  $[\zeta_o^2]_t \sim (\zeta_i - \zeta_o)\zeta_o + S$ 

**Correlation** between the ice and ocean vorticity occurs when c<80% and significantly **reduces** upper- ocean **dissipation** 



## Enhanced heat fluxes over anticyclones



km

-30